

1 A Bayesian statistical model for deriving the predictive 2 distribution of hydroclimatic variables

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6 **Abstract** Recent publications have provided evidence that hydrological processes exhibit a
7 scaling behaviour, also known as the Hurst phenomenon. An appropriate way to model this
8 behaviour is to use the Hurst-Kolmogorov stochastic process. The Hurst-Kolmogorov process
9 entails high autocorrelations even for large lags, as well as high variability even at climatic
10 scales. A problem that, thus, arises is how to incorporate the observed past hydroclimatic data
11 in deriving the predictive distribution of hydroclimatic processes at climatic time scales. Here
12 with the use of Bayesian techniques we create a framework to solve the aforementioned
13 problem. We assume that there is no prior information for the parameters of the process and
14 use a non-informative prior distribution. We apply this method with real-world data to derive
15 the posterior distribution of the parameters and the posterior predictive distribution of various
16 30-year moving average climatic variables. The marginal distributions we examine are the
17 normal and the truncated normal (for nonnegative variables). We also compare the results
18 with two alternative models, one that assumes independence in time and one with Markovian
19 dependence, and the results are dramatically different. The conclusion is that this framework
20 is appropriate for the prediction of future hydroclimatic variables conditional on the
21 observations.

22 **Key words** Bayesian statistics; hydroclimatic prediction; likelihood; Hurst-Kolmogorov
23 process; hydrological statistics.

24 **1 Introduction**

25 A lot of work has been done in predicting the future of hydroclimatic processes using
26 Bayesian statistics. Berliner et al. (2000) applied a Markov model to a low-order dynamical
27 system of tropical Pacific SST, using a hierarchical Bayesian dynamical modelling, which led
28 to realistic error bounds on forecasts. Duan et al. (2007) illustrated how the Bayesian model
29 averaging (BMA) scheme can be used to generate probabilistic hydrologic predictions from
30 several competing individual predictions. Kumar and Maity (2008) used two different
31 Bayesian dynamic modelling approaches, namely a constant model and a dynamic regression
32 model (DRM) to forecast the volume of the Devil's lake. Maity and Kumar (2006) used a
33 Bayesian dynamic linear model to predict the monthly Indian summer monsoon rainfall.
34 Bakker and Hurk (2012) used a Bayesian model to predict multi-year geostrophic winds.

35 On the other hand, climate models (i.e. general circulation models—GCMs) give
36 deterministic projections of future hydroclimatic processes for some hypothesized scenarios
37 e.g. for the increase of CO₂ concentration, etc. However, the uncertainty of these projections
38 whose sources may be attributed to insufficient current understanding of climatic
39 mechanisms, to inevitable weaknesses of numerical climatic and hydrologic models to
40 represent processes and scales of interest, to complexity of processes and to unpredictability
41 of causes (Koutsoyiannis et al. 2007), is not estimated by these models. Consequently, it is
42 impossible to estimate whether any observed changes reflect the natural variability of the
43 climatic processes or should be attributed to external forcings. Additionally, using
44 deterministic projections and thus neglecting the uncertainty in future hydroclimatic
45 conditions, may result in underestimation of possible range of the future hydroclimatic
46 variation.

47 Koutsoyiannis et al. (2007) have done some work on the uncertainty assessment of future
48 hydroclimatic predictions. They propose a stochastic framework for future climatic

49 uncertainty, where climate is expressed by the 30-year time average of a natural process
50 exhibiting a scaling behaviour, also known as the Hurst phenomenon or Hurst-Kolmogorov
51 (HK) behaviour (Hurst 1951; Koutsoyiannis et al. 2008). To this end, they combine analytical
52 and Monte Carlo methods to determine uncertainty limits and they apply the framework
53 developed to temperature, rainfall and runoff data from a catchment in Greece, for which
54 measurements are available for about a century.

55 In the study by Koutsoyiannis et al. (2007), the climatic variability and the influence of
56 parameter uncertainty are studied separately. As a result, a hydroclimatic prediction needs two
57 confidence coefficients to be defined, one referring to the uncertainty of the climatic evolution
58 and one to the uncertainty of model parameters. In this paper we unify the study of the two
59 uncertainties so that a climatic prediction needs only one confidence coefficient to be defined.
60 To this end, we solve the problem of climatic predictions of natural processes using Bayesian
61 statistics, instead of the stochastic framework developed by Koutsoyiannis et al. (2007). For
62 physical consistency with natural processes such as rainfall and runoff, whose values are
63 nonnegative, we also examine the case where truncation of the negative part of the
64 distributions is applied. No prior information for the parameters of processes is assumed, so
65 that the prior distribution is non-informative. The posterior joint distribution is derived from a
66 mixture for the case where truncation is not applied and a Gibbs sampler for the case where
67 truncation is applied. We derive the posterior predictive distribution (Gelman et al. 2004, p.8)
68 of the process in closed form given the posterior distribution of the parameters. We simulate a
69 sample from the posterior predictive distribution and use it to make inference about the future
70 evolution of the averaged process. We apply this procedure using the same data as in
71 Koutsoyiannis et al. (2007), and specifically runoff (Case 1 or C1), rainfall (C2) and
72 temperature (C3) data from catchments in Greece and temperature data from Berlin (C4, C6
73 with the last 90 years excluded from the dataset); in addition we used temperature data from

74 Vienna (C5, C7 with the last 90 years excluded from the dataset). For the rainfall and runoff
75 data we use truncated distributions.

76 As per the temporal dependence of the processes, three alternative assumptions are made:
77 (a) independence in time; (b) Markovian dependence modelled by first-order autoregressive
78 (AR(1)) process; and (c) HK dependence (see Markonis and Koutsoyiannis 2013, for a
79 justification of the latter). In the last section we compare the results of the three models.
80 Additional results such as the posterior distributions of the parameters and the asymptotic
81 behaviour of the predictive distribution are also given.

82 While this paper uses the same case studies as those in Koutsoyiannis et al. (2007), the
83 results of the two papers are not directly comparable to each other. Here we give posterior
84 predictive distributions of the climatic variables, whereas Koutsoyiannis et al. (2007) give
85 confidence limits for specified quantiles of climatic variables. The posterior predictive
86 distribution of the variables given here is exactly what we call climatic prediction, whereas we
87 could say that the confidence limits of the quantiles, given by Koutsoyiannis et al. (2007), are
88 intermediate or indirect results. The Bayesian methodology applied here aims at (stochastic)
89 prediction (Robert 2007, p.7) and is direct, while its disadvantage compared to Koutsoyiannis
90 et al. (2007) framework is the much heavier computational burden.

91 **2 Definition of AR(1) and HK process**

92 We use the Dutch convention for notation, according to which random variables and
93 stochastic processes are underlined (Hemelrijk 1966). We assume that $\{\underline{x}_t\}$, $t = 1, 2, \dots$ is a
94 normal stationary stochastic process with mean $\mu := E[\underline{x}_t]$, standard deviation $\sigma := \sqrt{\text{Var}[\underline{x}_t]}$,
95 autocovariance function $\gamma_k := \text{Cov}[\underline{x}_t, \underline{x}_{t+k}]$ and autocorrelation function (ACF) $\rho_k := \text{Corr}[\underline{x}_t, \underline{x}_{t+k}] = \gamma_k / \gamma_0$ ($k = 0, \pm 1, \pm 2, \dots$) and that there is a record of n observations $\mathbf{x}_n = (x_1 \dots x_n)^T$.
96 Each observation x_t represents a realization of a random variable \underline{x}_t , so that \mathbf{x}_n is a realization

98 of the vector of random variables $\underline{\mathbf{x}}_n = (\underline{x}_1 \dots \underline{x}_n)^T$.

99 We assume that $\{\underline{a}_t\}$ is a zero mean normal white noise process (WN), i.e. a sequence of
 100 independent random variables from a normal distribution with mean $E[\underline{a}_t] = 0$ and variance
 101 $\text{Var}[\underline{a}_t] = \sigma_a^2$. In the following discussion $\{\underline{a}_t\}$ is always referred to as WN. The following
 102 equation defines the first-order autoregressive process AR(1).

$$103 \quad \underline{x}_t - \mu = \varphi_1(\underline{x}_{t-1} - \mu) + \underline{a}_t, |\varphi_1| < 1 \quad (1)$$

104 The ACF of the AR(1) is (Wei 2006, p.34)

$$105 \quad \rho_k = \varphi_1^k, k = 0, 1, \dots \quad (2)$$

106 Let κ be a positive integer that represents a timescale larger than 1, the original time scale
 107 of the process \underline{x}_t . The averaged stochastic process on that timescale is denoted as

$$108 \quad \underline{x}_t^{(\kappa)} := (1/\kappa) \sum_{l=(t-1)\kappa+1}^{t\kappa} \underline{x}_l \quad (3)$$

109 The notation implies that a superscript (1) could be omitted, i.e. $\underline{x}_t^{(1)} \equiv \underline{x}_t$. Now we consider
 110 the following equation that defines the Hurst-Kolmogorov stochastic process (HKp).
 111 (Koutsoyiannis 2003)

$$112 \quad (\underline{x}_i^{(\kappa)} - \mu) \stackrel{d}{=} \left(\frac{\kappa}{\lambda}\right)^{H-1} (\underline{x}_j^{(\lambda)} - \mu), 0 < H < 1, i, j = 1, 2, \dots \text{ and } \kappa, \lambda \geq 1 \quad (4)$$

113 where H is the Hurst parameter.

114 The ACF of the HKp is (Koutsoyiannis 2003)

$$115 \quad \rho_k = |k+1|^{2H}/2 + |k-1|^{2H}/2 - |k|^{2H}, k = 0, 1, \dots \quad (5)$$

116 and does not depend on averaging time scale κ .

117 3 Posterior distribution of the parameters of a stationary normal stochastic 118 process

119 The distribution of the variable $\underline{\mathbf{x}}_n = (x_1 \dots x_n)^T$ is

$$120 \quad f(\mathbf{x}_n|\boldsymbol{\theta}) = (2\pi)^{-n/2} |\sigma^2 \mathbf{R}_n|^{-1/2} \exp[(-1/2\sigma^2) (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n)] \quad (6)$$

121 where \mathbf{R}_n is the autocorrelation matrix with elements $r_{ij} = \rho_{|i-j|}$, $i, j = 1, 2, \dots, n$ and $\mathbf{e}_n = (1 \ 1 \ \dots$
122 $1)^T$ is a vector with n elements. Details on the distributions used thereafter are given in
123 Appendix 1. The autocorrelation $\rho_{|i-j|}$ is assumed to be function of a parameter (scalar or
124 vector) $\boldsymbol{\varphi}$, so that $\boldsymbol{\theta} := (\mu, \sigma^2, \boldsymbol{\varphi})$ is the parameter vector of the process. We note that if $\underline{\mathbf{x}}_n$ is
125 white noise then $\rho_0 = 1$ and $\rho_k = 0$, $k = 1, 2, \dots$; if it is AR(1) then ρ_k is given by (2) if it is
126 HKp then ρ_k is given by (5).

127 We assume that $\boldsymbol{\varphi}$ is uniformly distributed a priori. We set as prior distribution for $\boldsymbol{\theta}$ the
128 non-informative distribution (see also Robert 2007, example 3.5.6)

$$129 \quad \pi(\boldsymbol{\theta}) \propto 1/\sigma^2 \quad (7)$$

130 (notice that we generally use the symbol π for probability density functions of parameters).

131 The posterior distribution of the parameters does not have a closed form. However it can
132 be calculated from a mixture based on conditional distributions. Specifically, it is shown (see
133 Appendix 2) that

$$134 \quad \underline{\mu}|\sigma^2, \boldsymbol{\varphi}, \mathbf{x}_n \sim N[(\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)/(\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n), \sigma^2/(\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)] \quad (8)$$

$$135 \quad \sigma^2|\boldsymbol{\varphi}, \mathbf{x}_n \sim \text{Inv-gamma}\{(n-1)/2, [\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)^2]/(2 \mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)\} \quad (9)$$

$$136 \quad \pi(\boldsymbol{\varphi}|\mathbf{x}_n) \propto |\mathbf{R}_n|^{-1/2} [\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)^2]^{-(n-1)/2} (\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)^{n/2-1} \quad (10)$$

137 As real world problems often impose upper or lower bounds on the variables $\underline{\mathbf{x}}_n$, we assume
138 that the distribution of $\underline{\mathbf{x}}_n$ is two-sided truncated by bounds a and b , i.e.,

$$f(\mathbf{x}_n|\boldsymbol{\theta}) \propto \exp\left[-\frac{1}{2\sigma^2} (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n)\right] I_{[a,b]^n}(x_1, \dots, x_n) \quad (11)$$

where I denotes the indicator function, so that $I_{[a,b]^n}(x_1, \dots, x_n) = 1$ if $\mathbf{x}_n \in [a,b]^n$ and 0 otherwise.

We assume that the truncation set of μ is $[a,b]$, $a, b \in \mathbb{R} \cup \{-\infty, \infty\}$. The following Gibbs sampler is used to obtain a posterior sample from $\underline{\boldsymbol{\theta}} = (\underline{\mu}, \underline{\sigma}^2, \underline{\boldsymbol{\varphi}})$ (see Appendix 2).

$$\pi(\mu|\sigma^2, \boldsymbol{\varphi}, \mathbf{x}_n) \propto \exp\left\{-\left[\mu - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n) / (\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)\right]^2 / (2\sigma^2 / \mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)\right\} I_{[a,b]}(\mu) \quad (12)$$

$$\underline{\sigma}^2 | \mu, \boldsymbol{\varphi}, \mathbf{x}_n \sim \text{Inv-gamma}\{n/2, (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n) / 2\} \quad (13)$$

$$\pi(\boldsymbol{\varphi} | \mu, \sigma^2, \mathbf{x}_n) \propto |\mathbf{R}_n|^{-1/2} \exp\left[-(\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n) / 2\sigma^2\right] \quad (14)$$

4 Posterior predictive distributions

As we stated in the Introduction, we seek to make an inference about the future evolution of a process given observations of its past. To this end, in this section we derive the posterior predictive distributions of $\underline{\mathbf{x}}_{n+1, n+m} | \mathbf{x}_n$ for the cases of the white noise, the AR(1) and the HKp, where $\underline{\mathbf{x}}_{n+1, n+m} := (\underline{x}_{n+1}, \dots, \underline{x}_{n+m})^T$.

4.1 White noise

We assume that \underline{x}_t , $t = 1, 2, \dots$ is white noise, with $f(x_t | \mu, \sigma^2) = (2\pi\sigma^2)^{-1/2} \exp[-(x_t - \mu)^2 / (2\sigma^2)]$. A non-informative prior distribution for $\underline{\boldsymbol{\theta}} = (\underline{\mu}, \underline{\sigma}^2)$ is $\pi(\boldsymbol{\theta}) \propto 1/\sigma^2$. The posterior distributions of the parameters are given by (Gelman et al. 2004, p.75-77)

$$\underline{\mu} | \mathbf{x}_n \sim t_{n-1}(\bar{x}_n, s_n^2 / (n-1)) \quad (15)$$

$$\underline{\sigma}^2 | \mathbf{x}_n \sim \text{Inv-gamma}((n-1)/2, n s_n^2 / 2) \quad (16)$$

Notice that (15) and (16) are derived from (8),(9),(10) for $\mathbf{R}_n = \mathbf{I}_n$ (the former after integrating

159 out σ^2). The posterior predictive distribution is

$$160 \quad \underline{x}_t | \mathbf{x}_n \sim t_{n-1}(\bar{x}_n, ((n+1)/(n-1))s_n^2), t = n+1, n+2, \dots \quad (17)$$

161 where $\underline{x}_{n+1}, \underline{x}_{n+2}, \dots$ are mutually independent,

$$162 \quad \bar{x}_n := \sum_{i=1}^n x_i/n \quad (18)$$

$$163 \quad s_n^2 := \sum_{i=1}^n (x_i - \bar{x}_n)^2/n \quad (19)$$

164 are the maximum likelihood estimates of μ and σ^2 respectively and $t_\nu(\mu, \sigma^2)$ is the Student's
165 distribution with ν degrees of freedom.

166 4.2 AR(1) and HKp

167 When there is dependence among the elements of $\underline{\mathbf{x}}_{n+m}$, the posterior predictive distribution of
168 $\underline{\mathbf{x}}_{n+1, n+m}$ given $\boldsymbol{\theta}$ and \mathbf{x}_n is (Eaton 1983, p.116,117)

$$169 \quad f(\underline{\mathbf{x}}_{n+1, n+m} | \boldsymbol{\theta}, \mathbf{x}_n) = (2\pi\sigma^2)^{-m/2} |\mathbf{R}_{m|n}|^{-1/2} \exp\left[(-1/2\sigma^2) (\underline{\mathbf{x}}_{n+1, n+m} - \boldsymbol{\mu}_{m|n})^T \mathbf{R}_{m|n}^{-1} (\underline{\mathbf{x}}_{n+1, n+m} - \boldsymbol{\mu}_{m|n})\right] \quad (20)$$

170 where $\boldsymbol{\mu}_{m|n}$ and $\mathbf{R}_{m|n}$ are given by:

$$171 \quad \boldsymbol{\mu}_{m|n} = \boldsymbol{\mu} \mathbf{e}_m + \mathbf{R}_{[(n+1):(n+m)] [1:n]} \mathbf{R}_n^{-1} (\mathbf{x}_n - \boldsymbol{\mu} \mathbf{e}_n) \quad (21)$$

$$172 \quad \mathbf{R}_{m|n} = \mathbf{R}_{[(n+1):(n+m)] [(n+1):(n+m)]} - \mathbf{R}_{[1:n] [(n+1):(n+m)]}^T \mathbf{R}_n^{-1} \mathbf{R}_{[1:n] [(n+1):(n+m)]} \quad (22)$$

173 where $\mathbf{R}_{[k:l] [m:n]}$ is the submatrix of \mathbf{R} which contains the elements r_{ij} , $k \leq i \leq l$, $m \leq j \leq n$,
174 whereas the notation $\mathbf{R}_{[1:n] [1:n]}$ with identical subscripts $[1:n]$ can be simplified to \mathbf{R}_n as
175 defined above. The elements of the correlation matrices \mathbf{R}_n and \mathbf{R}_{m+n} are obtained from (2) for
176 the case of the AR(1) and from (5) for the case of HKp. In the implementation of the AR(1)
177 model we assume that all three parameters μ , σ , φ_1 are unknown. For the HKp we examine

178 two cases: (a) all three parameters μ , σ , H , are unknown, and (b) μ , σ , are unknown but H is
 179 considered to be known and equal to its maximum likelihood estimate (Tyralis and
 180 Koutsoyiannis 2011).

181 In the case that all three parameters of the AR(1) or HKp are unknown, we obtain a
 182 simulated sample of $\underline{\theta}$ from (8),(9),(10) and use this sample to simulate $\underline{\mu}_{m|n}$ and $\underline{\mathbf{R}}_{m|n}$ from
 183 (21) and (22) and generate a sample of $\underline{\mathbf{x}}_{n+1,n+m}$ from (20). In the case where H is considered as
 184 known, we obtain a simulated sample of $\underline{\theta} = (\underline{\mu}, \underline{\sigma}^2)$ from (8),(9) and use this sample to
 185 simulate $\underline{\mu}_{m|n}$ and $\underline{\mathbf{R}}_{m|n}$ from (21) and (22) and generate a sample of $\underline{\mathbf{x}}_{n+1,n+m}$ from (20).

186 4.3 Asymptotic behaviour of AR(1) and HKp

187 In most applications, it is useful to know the ultimate confidence regions as prediction
 188 horizon tends to infinity. This is expressed by the distribution of $\underline{\mathbf{x}}_{n+m+1,n+m+l} :=$
 189 $(\underline{x}_{n+m+1}, \dots, \underline{x}_{n+m+l})$ as $m \rightarrow \infty$, conditional on \mathbf{x}_n . For given θ this distribution is:

$$190 f(\mathbf{x}_{n+m+1,n+m+l} | \theta, \mathbf{x}_n) = (2\pi\sigma^2)^{-l/2} |\mathbf{R}_{l|n}|^{-1/2} \exp\left[(-1/2\sigma^2)(\mathbf{x}_{n+m+1,n+m+l} - \boldsymbol{\mu}_{l|n})^T \mathbf{R}_{l|n}^{-1} (\mathbf{x}_{n+m+1,n+m+l} - \boldsymbol{\mu}_{l|n})\right] \quad (23)$$

191 where $\boldsymbol{\mu}_{l|n}$ and $\mathbf{R}_{l|n}$ are given by:

$$192 \boldsymbol{\mu}_{l|n} = \boldsymbol{\mu}e_l + \mathbf{R}_{[(n+m+1):(n+m+l)] [1:n]} \mathbf{R}_n^{-1} (\mathbf{x}_n - \boldsymbol{\mu}e_n) \quad (24)$$

$$193 \mathbf{R}_{l|n} = \mathbf{R}_{[(n+m+1):(n+m+l)] [(n+m+1):(n+m+l)]} - \mathbf{R}_{[1:n] [(n+m+1):(n+m+l)]}^T \mathbf{R}_n^{-1} \mathbf{R}_{[1:n] [(n+m+1):(n+m+l)]} \quad (25)$$

194 We observe that, as $m \rightarrow \infty$, $\mathbf{R}_{[1:n] [(n+m+1):(n+m+l)]}$ and $\mathbf{R}_{[(n+m+1):(n+m+l)] [1:n]}$ become zero matrices
 195 and $\mathbf{R}_{[(n+m+1):(n+m+l)] [(n+m+1):(n+m+l)]} = \mathbf{R}_l$. This implies that:

$$196 \boldsymbol{\mu}_{l|n} = \boldsymbol{\mu}e_l \quad (26)$$

$$197 \mathbf{R}_{l|n} = \mathbf{R}_l \quad (27)$$

198 where \mathbf{R}_l is again obtained from (2) for the case of the AR(1) and from (5) for the case of
 199 HKp.

200 Accordingly, the application can proceed as follows. We obtain a simulated sample of $\underline{\theta}$

201 from (8),(9),(10) and use this sample to simulate $\underline{\mu}_{l|n}$ and $\underline{R}_{l|n}$ from (26) and (27) and generate
 202 a sample of $\underline{x}_{n+m+1,n+m+l}$ from (23) for a large m .

203 4.4 Truncated white noise, AR(1) and HKp

204 To examine real world problems which often impose upper or lower bounds on the variables
 205 x_t , we assume that the distribution of \underline{x}_n is two-sided truncated, and is given by (11). We
 206 obtain a posterior sample of $\underline{\theta}$ using the Gibbs sampler defined by (12), (13), (14). When φ is
 207 known, we obtain a posterior sample of $(\underline{\mu}, \underline{\sigma}^2)$ using the Gibbs sampler defined by (12) and
 208 (13). Then $\underline{x}_m|\underline{\theta}$ follows a truncated normal multivariate distribution and according to Horrace
 209 (2005) the conditional multivariate distributions of $\underline{x}_{n+1,n+m}|\underline{\theta}, \underline{x}_n$ are again truncated normal.
 210 As a result (20) still holds after slight modifications and (21), (22) are valid. The posterior
 211 predictive distribution of $\underline{x}_{n+1,n+m}|\underline{\theta}, \underline{x}_n$ is then a multivariate truncated normal distribution:

$$212 \quad f(\underline{x}_{n+1,n+m}|\underline{\theta}, \underline{x}_n) \propto \exp\left[-\frac{1}{2\sigma^2} (\underline{x}_{n+1,n+m} - \underline{\mu}_{m|n})^T \underline{R}_{m|n}^{-1} (\underline{x}_{n+1,n+m} - \underline{\mu}_{m|n})\right] I_{[a,b]}^m(\underline{x}_{n+1,n+m}) \quad (28)$$

213 Now for the case of white noise, (15), (16) and (17) are not valid. But from (21), (22) and
 214 for $\rho_0 = 1$ and $\rho_k = 0, k = 1, 2, \dots$, we obtain that $\underline{\mu}_{m|n} = \underline{\mu} \underline{e}_m$ and $\underline{R}_{m|n} = \underline{R}_m$.

215 When looking for the asymptotic behaviour of the process, (23) still holds after slight
 216 modifications, according to Horrace (2005). As a result, the distribution of $\underline{x}_{n+m+1,n+m+l}|\underline{\theta}, \underline{x}_n$ is
 217 truncated multivariate normal, while (26) and (27) remain valid:

$$218 \quad f(\underline{x}_{n+m+1,n+m+l}|\underline{\theta}, \underline{x}_n) \propto \\
 219 \quad \propto \exp\left[-\frac{1}{2\sigma^2} (\underline{x}_{n+m+1,n+m+l} - \underline{\mu}_{l|n})^T \underline{R}_{l|n}^{-1} (\underline{x}_{n+m+1,n+m+l} - \underline{\mu}_{l|n})\right] I_{[a,b]}^l(\underline{x}_{n+m+1,n+m+l}) \quad (29)$$

220 4.5 Asymptotic convergence of MCMC

221 To simulate from (10) we use a random walk Metropolis-Hastings algorithm with a normal
 222 instrumental (or proposal) distribution (Robert and Casella 2004, p.271). We implement the
 223 algorithm using the function MCMCmetrop1R of the R package ‘MCMCpack’ (Martin et al.,

224 2011). The variable ‘burnin’ in this package is given the value 0, whereas the other variables
 225 keep their default values.

226 There are a lot of methods to decide whether convergence can be assumed to hold for the
 227 generated sample (see Gamerman and Lopes 2006, p.157-169; Robert and Casella 2004,
 228 p.272-276). We use the methods of Heidelberger and Welch (1983) and Raftery and Lewis
 229 (1992). These methods are described by Smith (2007), whose notation we use here. We use
 230 the R package ‘coda’ (Plummer et al. 2011) to implement these methods. We assume that we
 231 have obtained a sample ψ_1, ψ_2, \dots of a scalar variable φ using the MCMC algorithm.

232 The diagnostic of Heidelberger's method provides an estimate of the number of samples
 233 that should be discarded as a burn-in sequence and a formal test for non-convergence. The
 234 null hypothesis of convergence to a stationary chain is based on Brownian bridge theory and

235 uses the Cramer-von-Mises test statistic $\int_0^1 B_n(t)^2 dt$, where

$$236 \quad B_n(t) = (T_{\lfloor nt \rfloor} - \lfloor nt \rfloor \bar{\psi}) / \sqrt{nS(0)} \quad (30)$$

$$237 \quad T_k = \sum_{j=1}^k \psi_j, \quad k = 1, 2, \dots \text{ and } T_0 = 0 \quad (31)$$

238 where $\lfloor x \rfloor$ denotes the floor of x (the greatest integer not greater than x) and $S(0)$ is the
 239 spectral density evaluated at frequency zero. In calculating the test statistic, the spectral
 240 density is estimated from the second half of the original chain. If the null hypothesis is
 241 rejected, then the first $0.1n$ of the samples are discarded and the test is reapplied to the
 242 resulting chain. This process is repeated until the test is either non-significant or 50% of the
 243 samples have been discarded, at which point the chain is declared to be non-stationary. For
 244 more details see Smith (2007).

245 The methods of Raftery and Lewis are designed to estimate the number of MCMC samples

246 needed when quantiles are the posterior summaries of interest. Their diagnostic is applicable
247 for the univariate analysis of a single parameter and chain. For instance, let us consider the
248 estimation of the following posterior probability of a model parameter θ :

$$249 \quad P(f(\theta) < a \mid \mathbf{x}) = q \quad (32)$$

250 where \mathbf{x} denotes the observed data. Raftery and Lewis sought to determine the number of
251 MCMC samples to generate and the number of samples to discard in order to estimate q to
252 within $\pm r$ with probability s . In practice, users specify the values of q , r and s to be used in
253 applying the diagnostic (For more details see Smith, 2007).

254 To simulate from (14) we use an accept-reject algorithm (Robert and Casella 2004, p.51-
255 53) with a uniform instrumental density. Simulation from (12) and (13) is trivial. We assess
256 the convergence of the chain simulated from (12), (13), (14) using the method of Gelman and
257 Rubin (1992; see also Gelman 1996; Gamerman and Lopes 2006, p.166-168). An indicator of
258 convergence is formed by the estimator of a potential scale reduction (PSR) that is always
259 larger than 1. Convergence can be evaluated by the proximity of PSR to 1. Gelman (1996)
260 suggested accepting convergence when the value of PSR is below 1.2.

261 **5 Case studies**

262 In this section we apply the methodology developed in the previous sections to five historical
263 datasets; three of them obtained from the Boeotikos Kephisos River basin, one from Berlin
264 and one from Vienna. The choice of these datasets was dictated by the fact that they have
265 been also studied in other works with similar objectives, i.e. Koutsoyiannis et al. (2007) and
266 Koutsoyiannis (2011), so that the interested reader can make some comparisons. We present
267 the results of the application of the methodology to the aforementioned datasets.

268 **5.1 Historical datasets**

269 The first case study is performed on an important catchment in Greece, which is part of the

270 water supply system of Athens and has a history, as regards hydraulic infrastructure and
271 management that extends backward at least 3500 years. This is the closed (i.e. without outlet
272 to the sea) basin of the Boeoticos Kephisos River (Figure 1), with an area of 1955.6 km²,
273 mostly formed over a karstic subsurface. Owing to its importance for irrigation and water
274 supply, data availability for the catchment extends for about 100 years (the longest dataset in
275 Greece) and modelling attempts with good performance have already been carried out on the
276 hydrosystem (Rozos et al. 2004).

277 The long-term dataset for the basin extends from 1908 to 2003 and comprises a flow
278 record at the river outlet at the Karditsa station (C1), rainfall observations in the raingage
279 Aliartos (C2) and a temperature record at the same station (C3); the station locations are
280 shown in Figure 1. Further details on the construction of these datasets are given by
281 Koutsoyiannis et al. (2007). The relatively long records have already made it possible to
282 identify the scaling behaviour of rainfall and runoff in this basin (Koutsoyiannis 2003), and
283 make the catchment ideal for a case study of uncertainty assessment.

284 The two other datasets which we use are the mean annual temperature record of
285 Berlin/Templehof and Vienna, two of the longest series of instrumental meteorological
286 observations. For further details on the Berlin mean annual temperature dataset see
287 Koutsoyiannis et al. (2007) and for the Vienna mean annual temperature dataset see
288 Koutsoyiannis (2011). We examine two cases. In the first case we assume that the update of
289 the prior information is done (C4, C5), using the whole dataset. In the second case the update
290 is done excluding the last 90 years of the datasets (C6, C7).

291 **5.2 Application of the method**

292 We classified the data into three classes, the first containing the data from the Boeoticos
293 Kephisos River basin (C1-C3), the second containing the data from Berlin and Vienna (C4,
294 C6) and the third containing again the data from Berlin and Vienna (C5, C7) but excluding

295 the last 90 years. In the third case the posterior results were compared to the actual 90 last
296 years.

297 First we calculated the maximum likelihood estimates of the parameters for all the
298 examined cases (WN, AR(1), HKp). The results are given in Tables 1a and 1b. Truncated
299 models were used for C1 and C2 datasets due to the relatively high estimated σ which
300 otherwise would result in negative values. Instead, when we examined the temperature
301 datasets (C3-C7), simulated values near the absolute zero never appeared, indicating a good
302 behaviour of the non-truncated model.

303 The procedure for the temperature datasets is described below. We used (15) and (16) to
304 generate a posterior sample from $\underline{\mu}$ and $\underline{\sigma}^2$ for the WN case. To simulate from (10) for the φ_1
305 and \underline{H} posterior distribution of the AR(1) and HK cases correspondingly, we used a random
306 walk Metropolis-Hastings algorithm. We simulated a single chain with 3 000 000 MCMC
307 samples. The Metropolis acceptance rates are given in Table 2. To decide whether
308 convergence has been achieved, we used the Heidelberger and Welch method (1983). We
309 tested four cases, the first case containing all the 3 000 000 samples, the second containing the
310 last 2 000 000 samples and so forth. The results are presented in Tables 3a and 3b, from
311 where we conclude that stationary chain hypothesis holds in every case. We also used the
312 methods of Raftery and Lewis (1992), to estimate the number of MCMC samples needed
313 when quantiles are the posterior summaries of interest. The minimum number of samples and
314 the burn-in period for the simulation is given in Tables 4a and 4b, where $q = 0.025, 0.500,$
315 0.975 are the quantiles to be estimated, $r = 0.005$ is the desired margin of error of the estimate
316 and $s = 0.95$ is the probability of obtaining an estimate in the interval $(q-r, q+r)$. We decided
317 to use the last 2 000 000 samples of the chains, to obtain the histograms of the posterior
318 distributions of the parameters φ_1 and \underline{H} . The simulation of $\underline{\mu}, \underline{\sigma}^2$ from (8) and (9) is then
319 trivial. Summarized results for the parameters of the AR(1) and HK cases respectively are

320 shown in Tables 5a and 5b.

321 From the simulated samples we obtained the posterior probability plots of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for
 322 the AR(1) and HK cases (Figures 2 and 3). The last 100 000 simulated samples of the
 323 parameters, described in the previous paragraph were used to obtain samples from the
 324 required posterior predictive probabilities. The samples from the posterior predictive
 325 probability of $\underline{x}_t | \mathbf{x}_n$, $t = n+1, n+2, \dots, n+90$ were used to obtain samples for the variable of
 326 interest $\underline{x}_t^{(30)}$ given by (33).

$$327 \quad \underline{x}_t^{(30)} := (1/30) \left(\sum_{l=t-29}^n x_l + \sum_{l=n+1}^t \underline{x}_l \right), t = n+1, \dots, n+29 \text{ and } \underline{x}_t^{(30)} := (1/30) \sum_{l=t-29}^t \underline{x}_l, t = n+30, n+31, \dots (33)$$

328 We examined the cases of WN, AR(1), asymptotic behaviour of AR(1), HK where H is
 329 considered to be known and has the value of the maximum likelihood estimate, HK when H is
 330 not known, and its asymptotic behaviour. Figures 4, 5a, 5b show the 0.025, 0.500 and 0.975
 331 quantiles of the posterior predictive distributions of $\underline{x}_t^{(30)} | \mathbf{x}_n$, $t = n+1, n+2, \dots, n+90$.

332 The procedure for C1 and C2 is described below. We simulated from (12), (13) and (14) to
 333 obtain a posterior sample from $\underline{\mu}$, $\underline{\sigma}^2$ and $\underline{\varphi}$ for all cases. We simulated 10 chains with each one
 334 having 300 000 MCMC samples. To decide whether convergence has been achieved, we used
 335 the Gelman and Rubin (1992) rule. In all cases PSR ≈ 1 which shows that the chains
 336 converged to the target distribution. We decided to use the last 200 000 samples of each
 337 chain, to obtain the histograms of the posterior distributions of the parameters $\underline{\varphi}_1$ and \underline{H} .
 338 Summarized results for the parameters of the AR(1) and HK cases respectively are shown in
 339 Table 5a.

340 From the simulated samples we obtained the posterior probability plots of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for
 341 the AR(1) and HK cases (Figure 2a, 2b). The last 10 000 simulated samples of the parameters
 342 of each chain, described in the previous paragraph are used to obtain samples from the
 343 required posterior predictive probabilities. The samples from the posterior predictive

344 probability of $\underline{x}_t|\mathbf{x}_n$, $t = n+1, n+2, \dots, n+90$ are used to obtain samples for the variable of
 345 interest $\underline{x}_t^{(30)}$ given by (33). We examined the cases of WN, AR(1), asymptotic behaviour of
 346 AR(1), HK where H is considered to be known and has the value of the maximum likelihood
 347 estimate, HK with unknown H and its asymptotic behaviour. Figure 4 shows the 0.025, 0.500
 348 and 0.975 quantiles of the posterior predictive distributions of $\underline{x}_t^{(30)}|\mathbf{x}_n$, $t = n+1, n+2, \dots, n+90$.

349 5.3 Results

350 A first important result of the proposed framework is that it provides good estimates of the
 351 model parameters without introducing any assumptions (i.e., using non-informative priors).
 352 While common statistical methods give point estimates of parameters, the Bayesian
 353 framework provides also interval estimates based on their posterior distributions. The
 354 estimated values of μ are given in Table 6. It turns out that irrespective of the method used
 355 (MLE or posterior medians) they are almost equal. When examining temperatures, HKp
 356 resulted in the largest $\hat{\mu}$ and AR(1) in the second largest. In C4 and C6, $\hat{\mu}$ was larger than in
 357 C5 and C7 respectively. From the density diagrams of the posterior distributions (Figures 2-3)
 358 it seems that the posterior distribution of $\underline{\mu}$ is wider when HKp is used. The posterior
 359 distribution of $\underline{\sigma}$ is also wider on the right (see the values of the 0.975 quantiles in Tables
 360 5a,5b) for the HKp. However the estimated values of σ are almost equal for the three used
 361 models (Tables 1a and 1b). The estimated φ_1 and H are given in Tables 1a and 1b. Their
 362 estimated values for C5 are considerably higher compared to C7, but their posterior
 363 distributions are narrower (Table 5b), probably because of the bigger sample size in the
 364 former case. Their posterior distributions are also narrower for C4 compared to C6.

365 The second result of the framework is the predictive distribution of the future evolution of
 366 the process of interest. The posterior predictive 0.95-confidence regions for the 30-year
 367 moving averages are given in Figures 4, 5a and 5b. For C1 the confidence region is not
 368 symmetric with respect to the estimated mean, owing to the lower truncation bound alongside

369 with the relatively big $\hat{\sigma}$. In contrast, there is a symmetry for C2 owing to the relatively small
370 $\hat{\sigma}$, which justifies our decision to use models without truncation in those cases where $\hat{\sigma}$ is even
371 smaller (compared to mean). For all cases, the widest confidence regions correspond to the
372 HKp (due to the existence of persistence), followed by the AR(1), while the narrowest
373 confidence regions appear for the WN. Of course the confidence regions for unknown H are
374 wider than in the case where H was considered to be known and equal to its maximum
375 likelihood estimate. In C5 and C7 the HKp seems to be the best model, because it captures
376 better than the others the observed values of the climate variable for the last 90 years based on
377 the observed values of the previous years. In C7 it seems that the HKp did not capture the
378 increase of temperature in last decades. But when we examine the full dataset (C5), the
379 behaviour in last 90 years does not appear extraordinary. For the asymptotic values in the
380 HKp, the 0.95-confidence region ranges at intervals of the order of 150 mm (C1), 220 mm
381 (C2), 1.6°C (C3), 1.9°C (C4), 1.4°C (C5) for the 30-year moving average. The corresponding
382 values for the case of the WN of the order of 50 mm (C1), 75 mm (C2), 0.5°C (C3), 0.6°C
383 (C4), 0.6°C (C5) are considerably smaller compared to the case of the HKp.

384 **6 Summary**

385 We developed a Bayesian statistical methodology to make hydroclimatic prognosis in terms
386 of estimating future confidence regions on the basis of a stationary normal stochastic process.
387 We applied this methodology to five cases, namely the runoff (C1), the rainfall (C2) and the
388 temperature (C3) at Boeotikos Kephisos river basin in Greece, as well as the temperature at
389 Berlin (C4, C6) and the temperature at Vienna (C5, C7). The Bayesian statistical model
390 consisted of a stationary normal process (or truncated stationary normal process for the runoff
391 and rainfall cases) with a non-informative prior distribution. Three kinds of stationary normal
392 processes were examined, namely WN, AR(1) and HKp. We derived the posterior

393 distributions of the parameters of the models, the posterior predictive distributions of the
394 variables of the process and the posterior predictive distribution of the 30-year moving
395 average which was the climatic variable of interest. The methodology can also be applied to
396 other structures of the ACF.

397 A first important conclusion is that for all the examined cases and for all the examined
398 processes their estimated means are almost equal as expected. However the posterior
399 distributions of the means are wider when using the HKp, due to the persistence of the
400 process, and even wider when all parameters of the process are assumed to be unknown. This
401 results in wider confidence regions for future climatic variables of the processes. Moreover
402 the confidence regions of truncated future variables are asymmetric. This asymmetry depends
403 on the variance of the examined process. However the posterior distributions of the means of
404 all processes were less asymmetric.

405 Another important conclusion is that the use of short-range dependence stochastic
406 processes is not suitable to model geophysical processes, because they underestimate
407 uncertainty. However stationary persistent stochastic processes are suitable to achieve this
408 purpose. In the examined cases they performed well and were able to explain the fluctuations
409 of the process.

410 One may claim that, when climate is to be predicted, an assumption of stationarity is not an
411 appropriate one as currently several climate models project a changing future climate.
412 Nonetheless, an assessment of future climate variability and uncertainty based on the
413 stationarity hypothesis is a necessary step in establishing a stochastic method, whose
414 generalization at a second step would enable incorporating nonstationary components. In
415 addition, without knowing the variability under stationary conditions, it would not be possible
416 to quantify the credibility of climate models and even their usefulness. Work on the
417 generalization of the methodology to incorporate deterministic predictions by climate models

418 is under way and its results will be reported in due course.

419 Appendix 1: Standard probability distributions

420 For easy reference, the details of the distribution functions used in this paper are summarized
421 in Table 7.

422

423 **Table 7.** Distributions used in the Bayesian framework

Distribution	Notation	Parameters	Density function
Normal	$\underline{x} \sim N(\mu, \sigma^2)$	location μ scale $\sigma > 0$	$f_N(x \mu, \sigma^2) = (2\pi\sigma^2)^{-1/2} \exp[-(1/2\sigma^2)(x - \mu)^2]$
Truncated normal	$\underline{x} \sim TN(\mu, \sigma^2, a, b)$	location μ scale $\sigma > 0$ a minimum value b maximum value	$f_{TN}(x \mu, \sigma^2, a, b) = [f_N((b - \mu)/\sigma) - f_N((a - \mu)/\sigma)]^{-1} (1/\sigma) f_N((x - \mu)/\sigma)$ $x \in [a, b], f_N(x) := f_N(x 0, 1^2)$
Multivariate normal	$\underline{x} \sim N(\boldsymbol{\mu}, \boldsymbol{\Sigma})$ (implicit dimension n)	location $\boldsymbol{\mu}$ symmetric, pos. definite $n \times n$ variance matrix $\boldsymbol{\Sigma}$	$f_{MN}(\mathbf{x} \boldsymbol{\mu}, \boldsymbol{\Sigma}) = (2\pi)^{-n/2} \boldsymbol{\Sigma} ^{-1/2} \exp[-(1/2)(\mathbf{x} - \boldsymbol{\mu})^T \boldsymbol{\Sigma}^{-1} (\mathbf{x} - \boldsymbol{\mu})]$
Inverse-gamma	$\underline{x} \sim \text{Inv-gamma}(\alpha, \beta)$	shape $\alpha > 0$ scale $\beta > 0$	$f_{IG}(x \alpha, \beta) = \beta^\alpha [\Gamma(\alpha)]^{-1} x^{-(\alpha+1)} \exp(-\beta/x), x > 0$
Student-t	$\underline{x} \sim t_n(\mu, \sigma^2)$	degrees of freedom n location μ scale $\sigma > 0$	Not needed in the manuscript

424 Appendix 2: Mathematical proofs

425 In Appendix 2 the proofs of (8),(9),(10),(12),(13),(14) are given. It is easily shown that

$$426 (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n) = \mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mu^2 - 2 \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mu + \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n \quad (34)$$

427 After completing the squares the above expression becomes:

$$428 \mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mu^2 - 2 \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mu + \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n = \mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n [\mu - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n) / (\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)]^2 + [\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n$$

$$429 \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)^2 / (\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)] \quad (35)$$

430 From (6) and (7) we obtain the following:

$$431 \pi(\boldsymbol{\theta}) f(\mathbf{x}_n|\boldsymbol{\theta}) \propto \sigma^{-(n+2)} |\mathbf{R}_n|^{-1/2} \exp[-(1/2\sigma^2)(\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n)] \quad (36)$$

432 From (34),(35) and (36) we obtain (8). After integration of (36) we obtain (37) which proves

433 (9):

434
$$\pi(\sigma^2|\boldsymbol{\varphi}, \mathbf{x}_n) \propto (\sigma^2)^{-(n+1)/2} |\mathbf{R}_n|^{-1/2} \exp\left[(-1/2\sigma^2)[\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n \mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{x}_n - (\mathbf{x}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)^2]/(\mathbf{e}_n^T \mathbf{R}_n^{-1} \mathbf{e}_n)\right] \quad (37)$$

435 After integration of (36) we obtain (38), which proves (10) after integration:

436
$$\pi(\boldsymbol{\varphi}|\mathbf{x}_n) \propto \int \int \sigma^{-(n+2)} |\mathbf{R}_n|^{-1/2} \exp\left[(-1/2\sigma^2) (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n)\right] d\mu d\sigma^2 \quad (38)$$

437 See also Falconer and Fernandez (2007) for some results.

438 Now for the case where truncation is applied we obtain from (7) and (11):

439
$$\pi(\boldsymbol{\theta}) f(\mathbf{x}_n|\boldsymbol{\theta}) \propto \sigma^{-(n+2)} |\mathbf{R}_n|^{-1/2} \exp\left[(-1/2\sigma^2) (\mathbf{x}_n - \mu \mathbf{e}_n)^T \mathbf{R}_n^{-1} (\mathbf{x}_n - \mu \mathbf{e}_n)\right] I_{[a,b]^n}(x_1, \dots, x_n) \quad (39)$$

440 Conditional on $\mu \in [a,b]$, $a, b \in \mathbb{R} \cup \{-\infty, \infty\}$ the derivation of (12), (13) and (14) from (39) is
441 then trivial.

442 **Acknowledgements:** The authors wish to thank the eponymous reviewer Dr. Federico
443 Lombardo and an anonymous reviewer for their encouraging and constructive comments
444 which helped to improve the quality of the manuscript significantly.

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522

523 **Table 1a.** Summarized results and maximum likelihood estimates for the cases of WN, AR(1)
 524 and HKp at Boeoticos Kephisos River basin.

	Boeoticos basin		
	Runoff (mm)	Rainfall (mm)	Temperature (°C)
Start year	1908	1908	1898
End year	2003	2003	2003
Size, n	96	96	106
WN			
$\hat{\mu}$	197.63	658.36	16.96
$\hat{\sigma}$	81.25	155.82	0.69
AR(1)			
$\hat{\mu}$	197.65	658.22	16.96
$\hat{\sigma}$	81.22	155.81	0.69
$\hat{\varphi}_1$	0.34	0.10	0.31
HK			
$\hat{\mu}$	195.11	657.38	16.97
$\hat{\sigma}$	80.47	155.00	0.70
\hat{H}	0.71	0.60	0.71

525 **Table 1b.** Summarized results and maximum likelihood estimates for the cases of WN, AR(1)
 526 and HKp at Berlin and Vienna.

	Berlin	Vienna	Berlin	Vienna
	Temperature (°C)	Temperature (°C)	Temperature (°C)	Temperature (°C)
Start year	1756	1775	1756	1775
End year	2009	2009	1919	1919
Size, n	254	235	164	145
WN				
$\hat{\mu}$	9.17	9.58	9.04	9.36
$\hat{\sigma}$	0.91	0.87	0.92	0.84
AR(1)				
$\hat{\mu}$	9.18	9.58	9.05	9.36
$\hat{\sigma}$	0.92	0.87	0.92	0.84
$\hat{\varphi}_1$	0.37	0.30	0.30	0.11
HK				
$\hat{\mu}$	9.27	9.64	9.10	9.37
$\hat{\sigma}$	0.91	0.86	0.92	0.84
\hat{H}	0.73	0.70	0.70	0.59

527 **Table 2.** Metropolis acceptance rate for the MCMC simulation of φ_1 and \underline{H} , respectively, at
 528 Boeoticos Kephisos River basin.

	Aliartos temperature	Berlin temperature (1756-2009)	Vienna temperature (1775-2009)	Berlin temperature (1756-1919)	Vienna temperature (1775-1919)
φ_1	0.70731	0.70603	0.70612	0.70649	0.70654
\underline{H}	0.706037	0.70551	0.70599	0.70601	0.70638

529 **Table 3a.** Heidelberger and Welch test, for significance level 0.05, at Boeoticos Kephisos
 530 River basin.

Parameter	Aliartos temperature							
	ϱ_1	\underline{H}						
Stationarity test	passed	passed	passed	passed	passed	passed	passed	passed
Start iteration	1	1	1	1	1	1	1	1
p -value	0.427	0.745	0.46	0.242	0.869	0.567	0.338	0.618

531 **Table 3b.** Heidelberger and Welch test, for significance level 0.05, at Berlin and Vienna.

Data start	Berlin temperature (1756-2009)				Vienna temperature (1775-2009)			
	1	1000000	2000000	2900000	1	1000000	2000000	2900000
Parameter	ϱ_1				ϱ_1			
Stationarity test	passed	passed	passed	passed	passed	passed	passed	passed
Start iteration	1	1	1	1	1	1	1	1
p -value	0.943	0.738	0.342	0.448	0.928	0.696	0.366	0.0761
Parameter	\underline{H}				\underline{H}			
Stationarity test	passed	passed	passed	passed	passed	passed	passed	passed
Start iteration	1	1	1	1	1	1	1	1
p -value	0.837	0.466	0.279	0.691	0.789	0.501	0.296	0.84

Parameter	Berlin temperature (1756-1919)				Vienna temperature (1775-1919)			
	ϱ_1				ϱ_1			
Stationarity test	passed	passed	passed	passed	passed	passed	passed	passed
Start iteration	1	1	1	1	1	1	1	1
p -value	0.94	0.589	0.376	0.425	0.777	0.55	0.308	0.592
Parameter	\underline{H}				\underline{H}			
Stationarity test	passed	passed	passed	passed	passed	passed	passed	passed
Start iteration	1	1	1	1	1	1	1	1
p -value	0.833	0.606	0.339	0.923	0.885	0.83	0.373	0.323

532 **Table 4a.** Raftery and Lewis test for the case of Boeoticos Kephisos River basin.

q	Aliartos temperature							
	Burn- Total	Lower Dependence			Burn- Total	Lower Dependence		
ϱ_1	in		bound	factor	in		bound	factor
0.025	21	31794	3746	8.49	\underline{H} 18	35784	4899	7.3
0.500	24	356752	38415	9.29	24	464024	50239	9.24
0.975	28	32298	3746	8.62	28	42161	4899	8.61

533 Note: q is the quantile to be estimated, $r = 0.005$ is the desired margin of error of the
 534 estimate, $s = 0.95$ the probability of obtaining an estimate in the interval $(q-r, q+r)$, $\text{eps} =$
 535 0.001 is the precision required for estimating time to convergence.

536 **Table 4b.** Raftery and Lewis test for the cases of Berlin and Vienna.

		Berlin temperature (1756-2009)				Vienna temperature (1775-2009)			
q	Burn-in	Total	Lower bound	Dependence factor	Burn-in	Total	Lower bound	Dependence factor	
ϱ_1	0.025	21	31416	3746	8.39	21	31612	3746	8.44
	0.500	24	356512	38415	9.28	21	322441	38415	8.39
	0.975	21	31731	3746	8.47	21	31745	3746	8.47
\underline{H}	0.025	18	27288	3746	7.28	18	35670	4899	7.28
	0.500	21	322777	38415	8.4	21	422975	50239	8.42
	0.975	28	32732	3746	8.74	28	42882	4899	8.75
		Berlin temperature (1756-1919)				Vienna temperature (1775-1919)			
ϱ_1	0.025	21	31780	3746	8.48	21	31780	3746	8.48
	0.500	24	356656	38415	9.28	21	323631	38415	8.42
	0.975	21	32193	3746	8.59	21	32137	3746	8.58
\underline{H}	0.025	18	27330	3746	7.3	18	27072	3746	7.23
	0.500	21	323330	38415	8.42	21	324177	38415	8.44
	0.975	18	32991	3746	8.81	27	39690	3746	10.6

537 Note: q is the quantile to be estimated, $r = 0.005$ is the desired margin of error of the
538 estimate, $s = 0.95$ the probability of obtaining an estimate in the interval $(q-r, q+r)$, $\text{eps} =$
539 0.001 is the precision required for estimating time to convergence.

540 **Table 5a.** Summary results for the parameters of the AR(1) and HK cases at Boeoticos
541 Kephisos River basin.

Case	Mean	Standard Deviation	Quantiles				
			2.5%	25%	50%	75%	97.5%
Boeoticos runoff							
AR(1)							
μ	197.7	12.69	172.5	189.4	197.7	205.9	222.8
σ	83.93	7.41	71.50	78.78	83.23	88.29	100.45
ϱ_1	0.35	0.10	0.16	0.28	0.35	0.42	0.55
HK							
μ	194.85	31.30	132	178.1	195	211.6	256.1
σ	86.51	12.35	71.19	79.15	84.40	91.06	114.22
\underline{H}	0.74	0.07	0.62	0.69	0.74	0.78	0.88
Aliartos rainfall							
AR(1)							
μ	658.18	18.57	621.5	646	658.2	670.4	694.7
σ	159.9	12.24	138.3	151.3	159.1	167.5	186.2
ϱ_1	0.11	0.10	-0.09	0.04	0.11	0.18	0.32
HK							
μ	657.09	31.98	592.5	638.4	657.3	676.1	720.4
σ	160.7	13.45	137.9	151.4	159.5	168.6	190.3
\underline{H}	0.62	0.06	0.51	0.58	0.62	0.66	0.75
Aliartos temperature							
AR(1)							
μ	16.96	0.10	16.76	16.89	16.96	17.02	17.15
σ	0.71	0.06	0.61	0.67	0.70	0.74	0.84
ϱ_1	0.33	0.10	0.14	0.26	0.33	0.39	0.52
HK							
μ	16.97	0.29	16.44	16.83	16.97	17.11	17.52
σ	0.75	0.13	0.62	0.68	0.73	0.79	0.99
\underline{H}	0.74	0.07	0.61	0.69	0.74	0.79	0.88

542 **Table 5b.** Summary results for the parameters of the AR(1) and HK cases respectively at
 543 Berlin and Vienna.

Case	Mean	Standard Deviation	2.5%	25%	Quantiles		
					50%	75%	97.5%
Berlin temperature (1756-2009)							
AR(1)							
$\underline{\mu}$	9.18	0.09	9.01	9.12	9.18	9.24	9.35
$\underline{\sigma}$	0.93	0.05	0.84	0.89	0.92	0.96	1.03
$\underline{\varrho}_1$	0.38	0.06	0.26	0.34	0.38	0.42	0.49
<u>HK</u>							
$\underline{\mu}$	9.28	0.25	8.80	9.13	9.27	9.43	9.79
$\underline{\sigma}$	0.94	0.06	0.83	0.89	0.93	0.97	1.08
\underline{H}	0.75	0.03	0.67	0.72	0.75	0.77	0.83
Vienna temperature (1775-2009)							
AR(1)							
$\underline{\mu}$	9.58	0.08	9.42	9.53	9.58	9.63	9.74
$\underline{\sigma}$	0.88	0.05	0.80	0.85	0.88	0.91	0.98
$\underline{\varrho}_1$	0.31	0.06	0.19	0.27	0.31	0.35	0.43
<u>HK</u>							
$\underline{\mu}$	9.64	0.19	9.27	9.52	9.64	9.76	10.03
$\underline{\sigma}$	0.88	0.05	0.79	0.84	0.87	0.91	0.99
\underline{H}	0.71	0.04	0.64	0.68	0.71	0.73	0.79
Berlin temperature (1756-1919)							
AR(1)							
$\underline{\mu}$	9.05	0.10	8.85	8.98	9.05	9.12	9.25
$\underline{\sigma}$	0.94	0.06	0.83	0.89	0.93	0.97	1.06
$\underline{\varrho}_1$	0.31	0.08	0.16	0.26	0.31	0.37	0.46
<u>HK</u>							
$\underline{\mu}$	9.11	0.26	8.60	8.95	9.10	9.26	9.64
$\underline{\sigma}$	0.96	0.08	0.83	0.90	0.95	1.00	1.14
\underline{H}	0.72	0.05	0.63	0.69	0.72	0.76	0.83
Vienna temperature (1775-1919)							
AR(1)							
$\underline{\mu}$	9.36	0.08	9.20	9.31	9.36	9.42	9.52
$\underline{\sigma}$	0.86	0.05	0.76	0.82	0.85	0.89	0.97
$\underline{\varrho}_1$	0.12	0.08	-0.04	0.06	0.12	0.18	0.29
<u>HK</u>							
$\underline{\mu}$	9.37	0.13	9.10	9.29	9.37	9.45	9.63
$\underline{\sigma}$	0.86	0.06	0.76	0.82	0.86	0.89	0.98
\underline{H}	0.61	0.05	0.51	0.57	0.61	0.64	0.72

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545 **Table 6.** Estimates of μ using various methods.

Examined case	Maximum likelihood estimate			50% quantile	
	WN	AR(1)	HKp	AR(1)	HK
Boeoticos runoff	197.63	197.65	195.11	197.7	195
Aliartos rainfall	658.36	658.22	657.38	658.2	657.3
Aliartos temperature	16.96	16.96	16.97	16.96	16.97
Berlin temperature (1756-2009)	9.17	9.18	9.27	9.18	9.28
Vienna temperature (1775-2009)	9.58	9.58	9.64	9.58	9.64
Berlin temperature (1756-1919)	9.04	9.05	9.10	9.05	9.11
Vienna temperature (1775-1919)	9.36	9.36	9.37	9.36	9.37

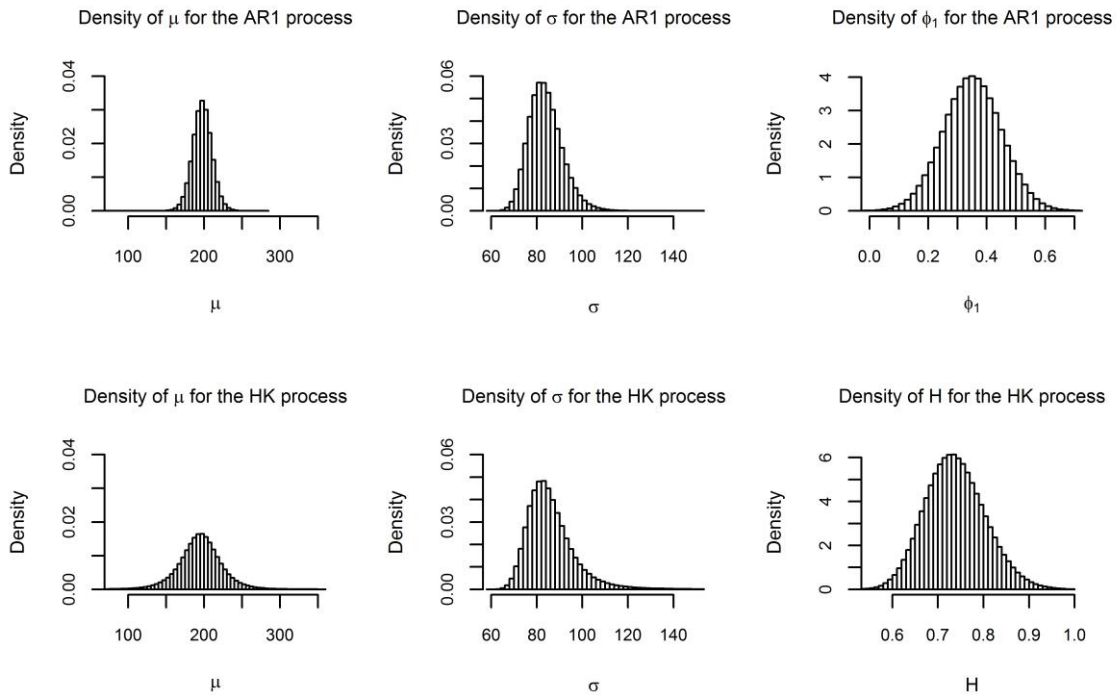
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Figure 1. The Boeoticos Kephisos River basin.

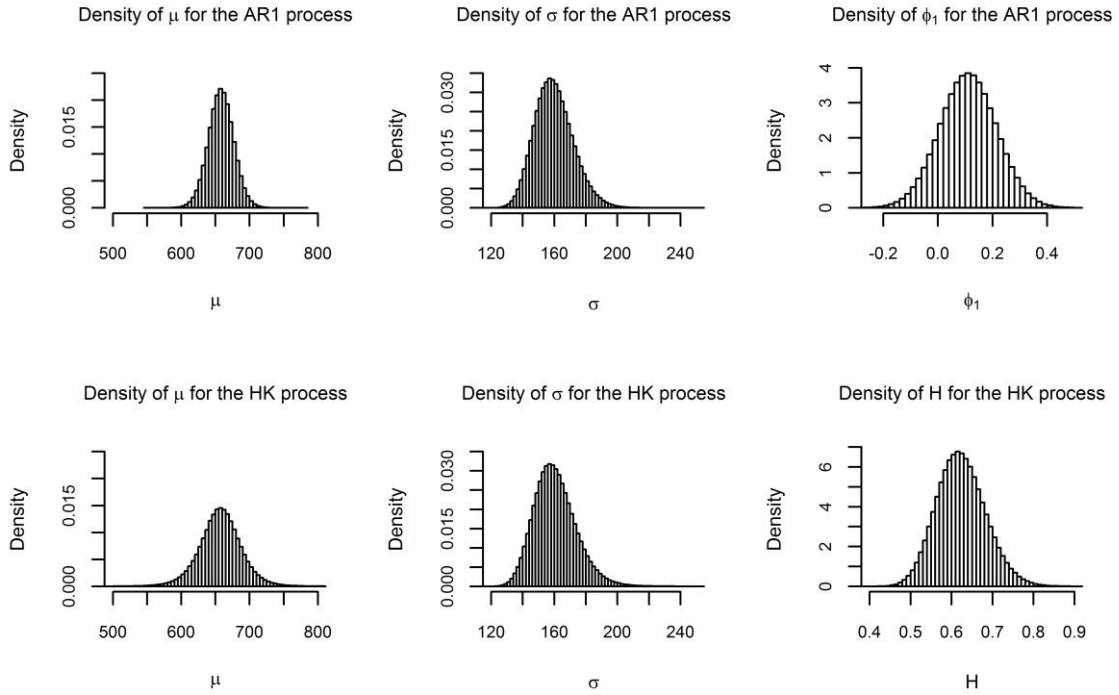


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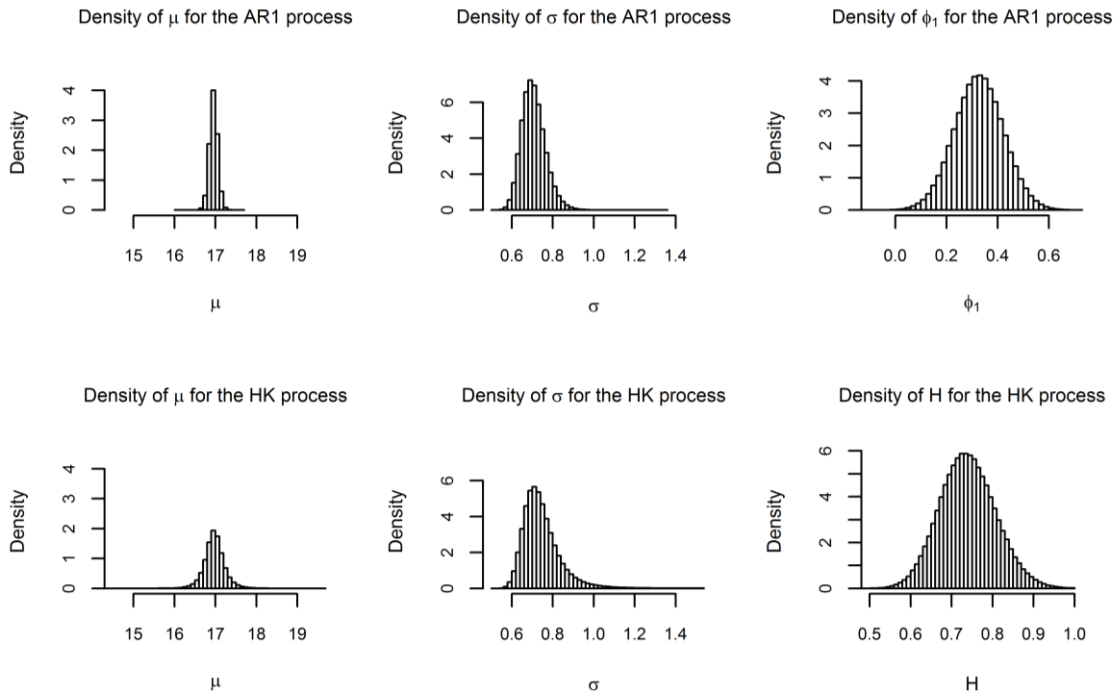
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Figure 2a. Posterior probability distributions of μ , σ , H , ϕ_1 for the cases of AR(1) and HK processes, for the runoff of Boeoticos Kephisos.



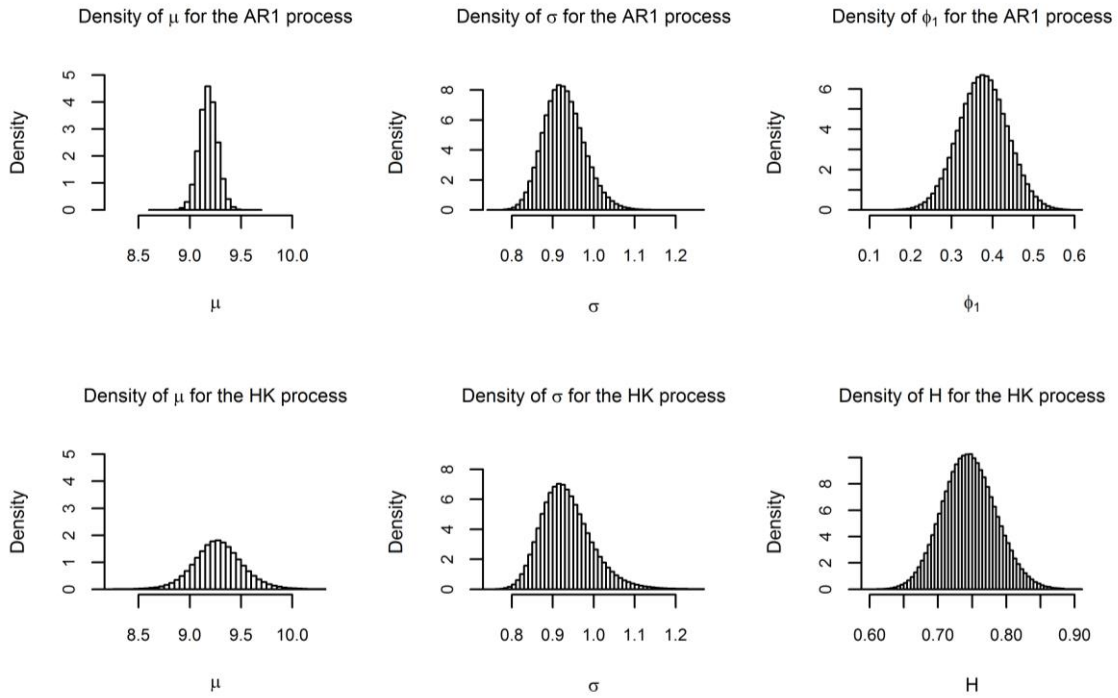
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Figure 2b. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the rainfall at Aliartos.



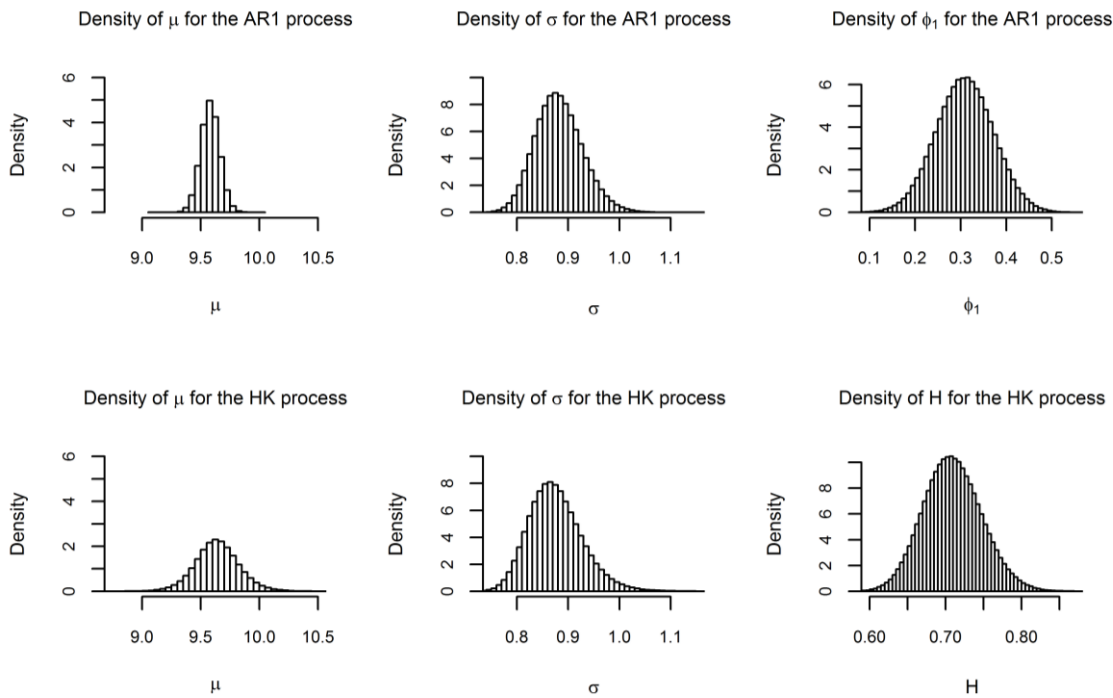
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Figure 2c. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the temperature at Aliartos.



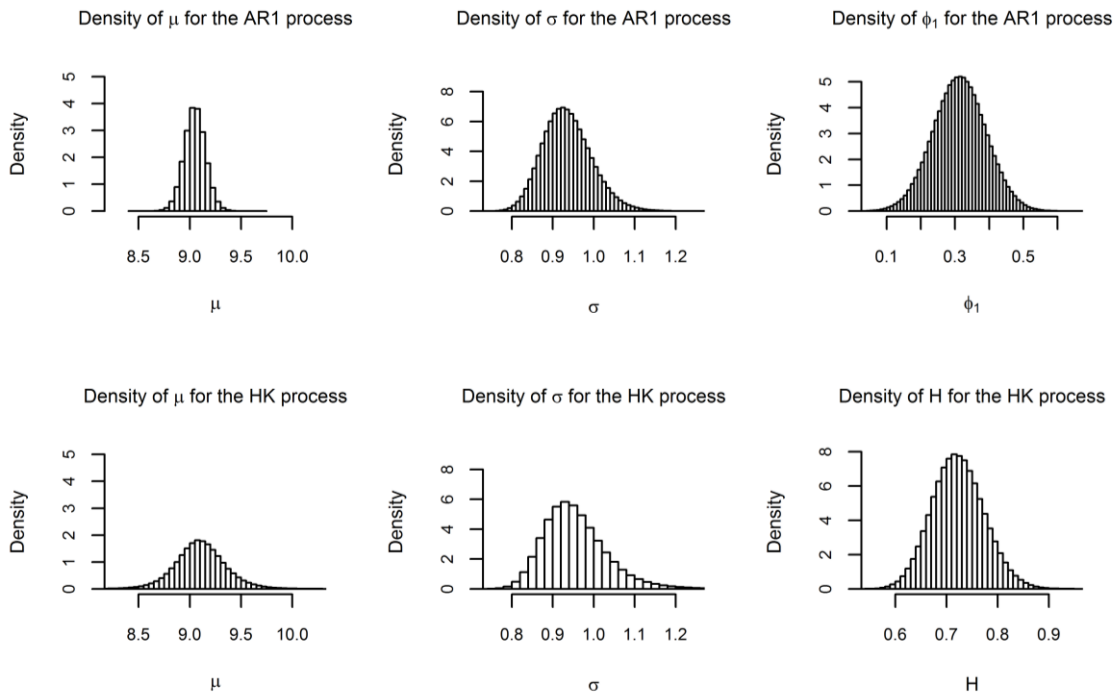
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Figure 3a. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the temperature at Berlin/Tempelhof. In this case the parameters are estimated from years 1756-2009.



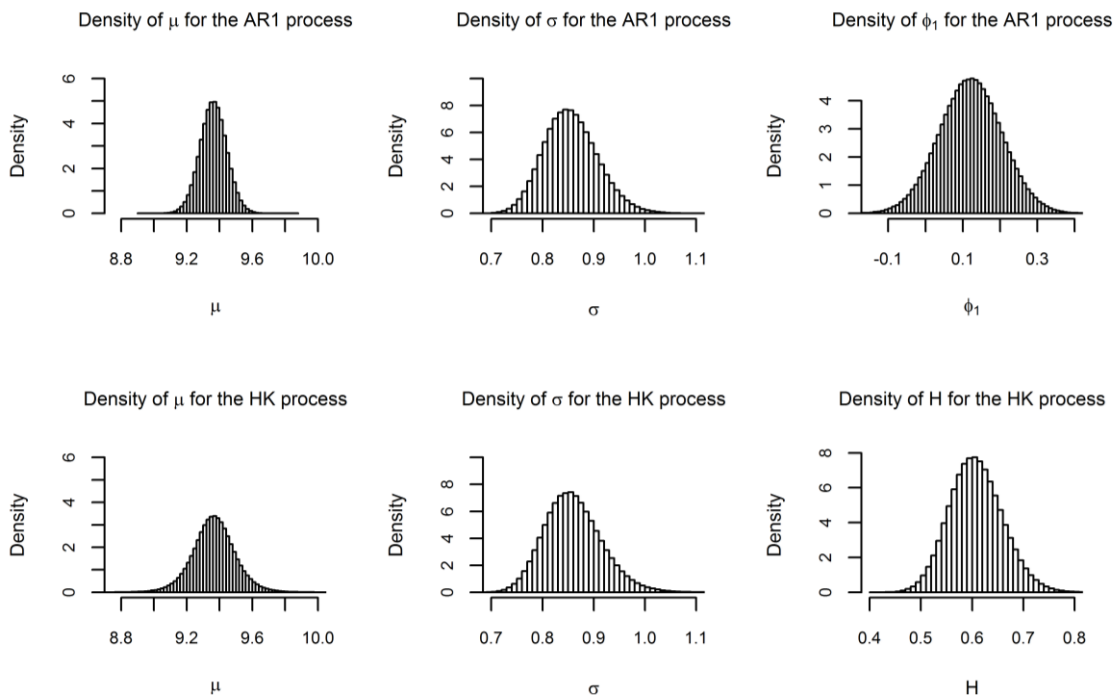
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Figure 3b. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the temperature at Vienna. In this case the parameters are estimated from years 1775-2009.



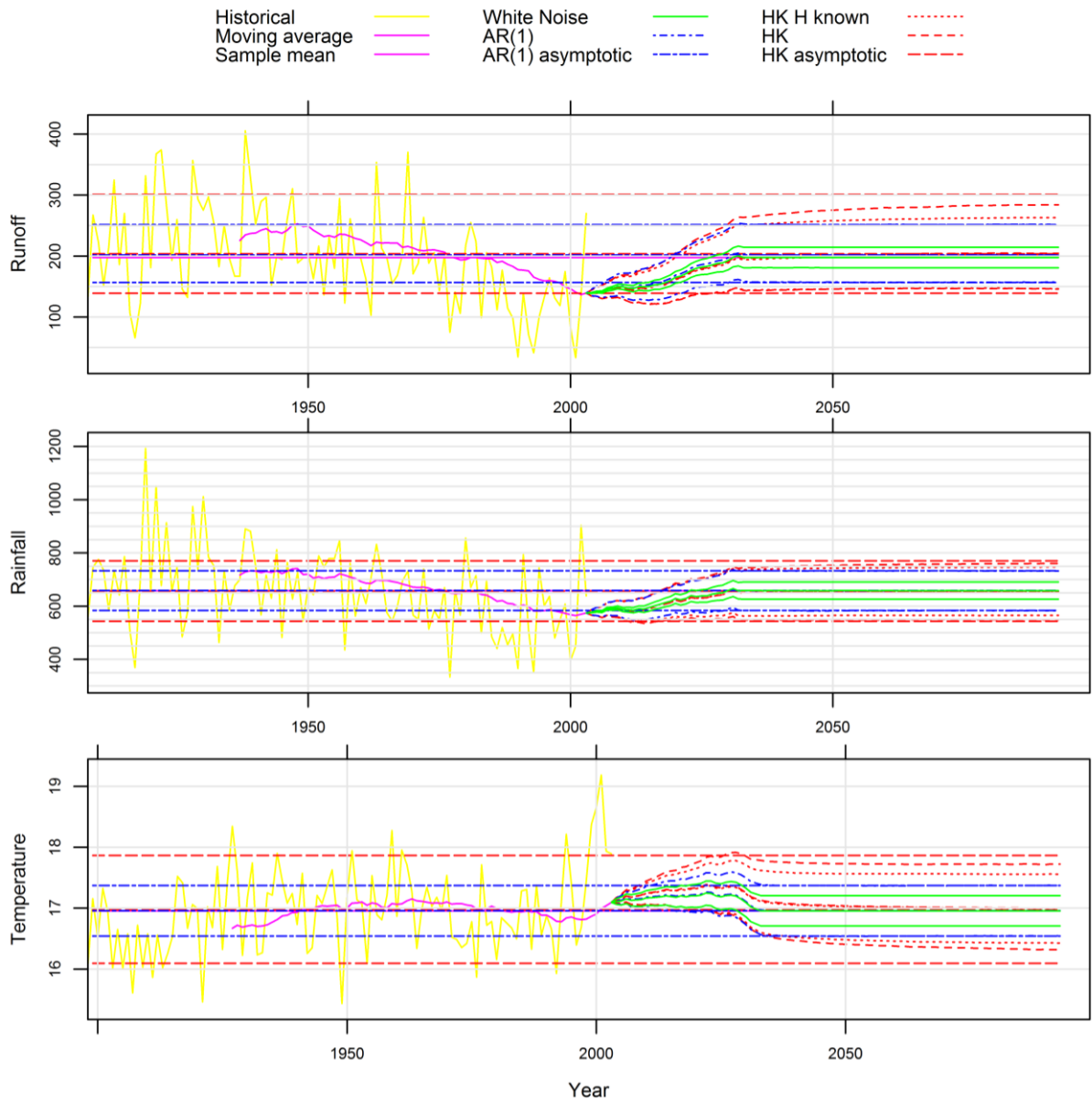
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Figure 3c. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the temperature at Berlin/Tempelhof. In this case the parameters are estimated from years 1756-1919.



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Figure 3d. Posterior probability distributions of $\underline{\mu}$, $\underline{\sigma}$, \underline{H} , $\underline{\varphi}_1$ for the cases of AR(1) and HK processes, for the temperature at Vienna. In this case the parameters are estimated from years 1775-1919.



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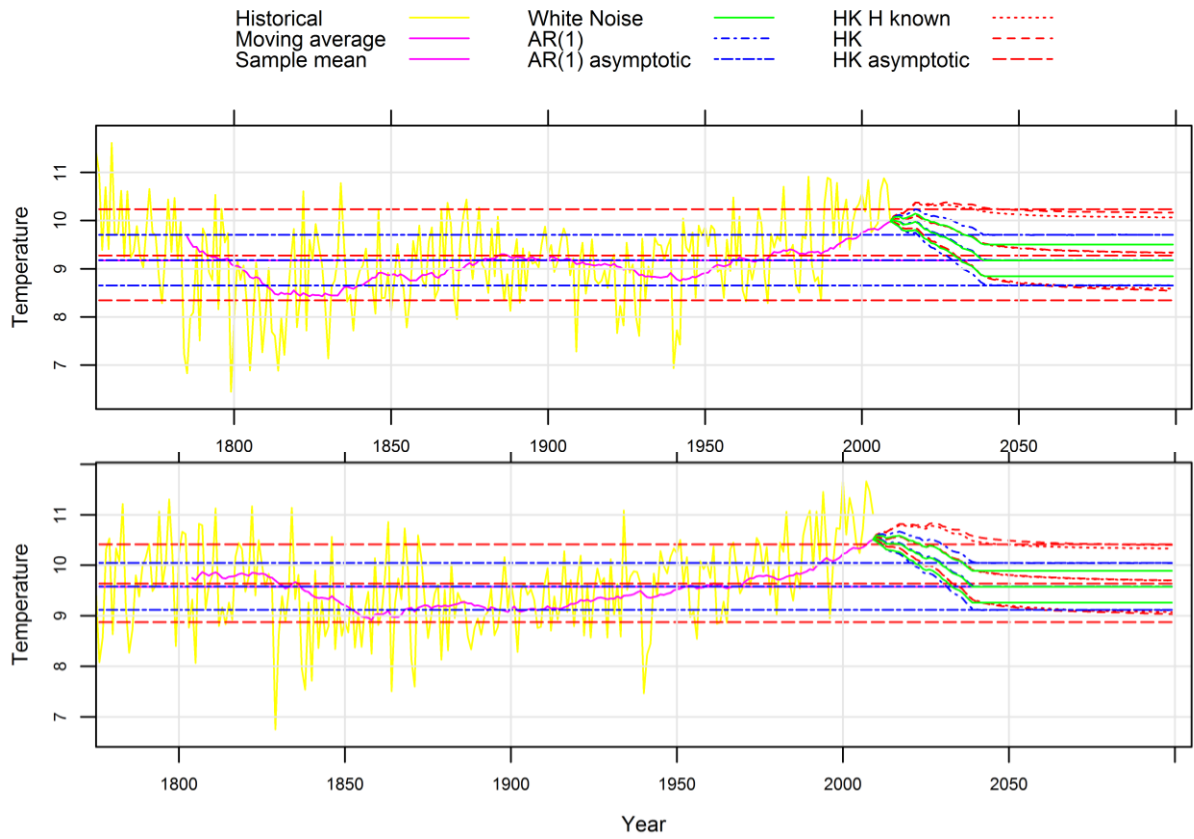
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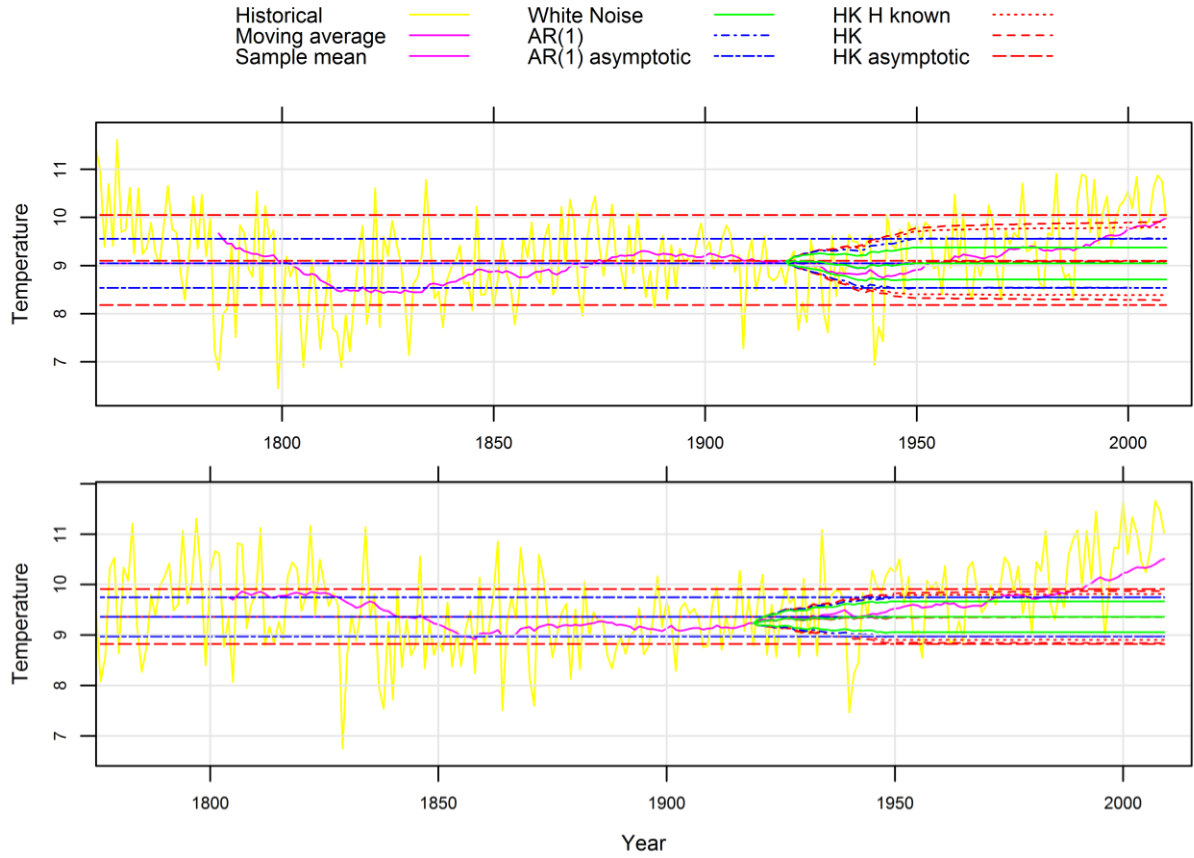
Figure 4. Historical climate and confidence regions of future climate (for $1 - a = 0.95$ and climatic time scale of 30 years) for (upper) runoff of Boeoticos Kephisos, (middle) rainfall at Aliartos, and (lower) temperature at Aliartos.



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Figure 5a. Historical climate and confidence regions of future climate (for $1 - \alpha = 0.95$ and climatic time scale of 30 years) for (upper) temperature at Berlin, and (lower) temperature at Vienna.



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Figure 5b. Historical climate and confidence regions of climate (for $1 - \alpha = 0.95$ and climatic time scale of 30 years) for (upper) temperature at Berlin/Tempelhof after the year 1920 and (lower) temperature at Vienna after the year 1920.