

*European Geosciences Union General Assembly
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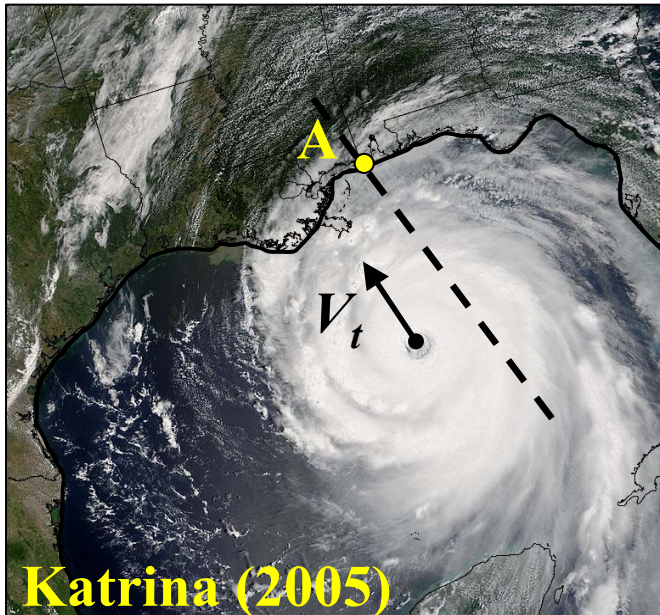
*Extreme Rainfall Intensities and
Long-term Rainfall Risk from
Tropical Cyclones*

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Objective

Long-term rainfall risk from TCs at location A:



$\lambda_D(i)$: rate at which $I_{max}(D)$ exceeds i at location A (events/year)

$I_{max}(D)$: maximum rainfall intensity at location A for averaging duration D

Risk analysis \Rightarrow

$$\lambda_D(i) = \lambda \int_{\text{all } \omega} P[I_{max}(D) > i | \omega] P[\omega] d\omega$$

λ : TC arrival rate [events/yr]

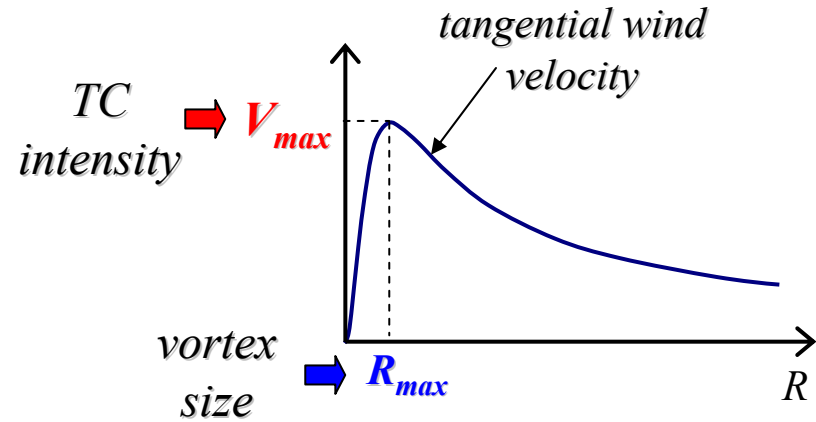
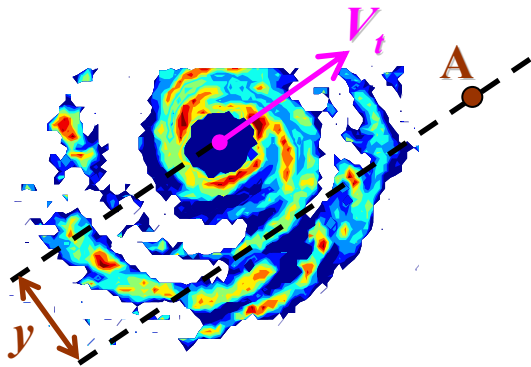
$P[\omega]$: TC characteristics

focus

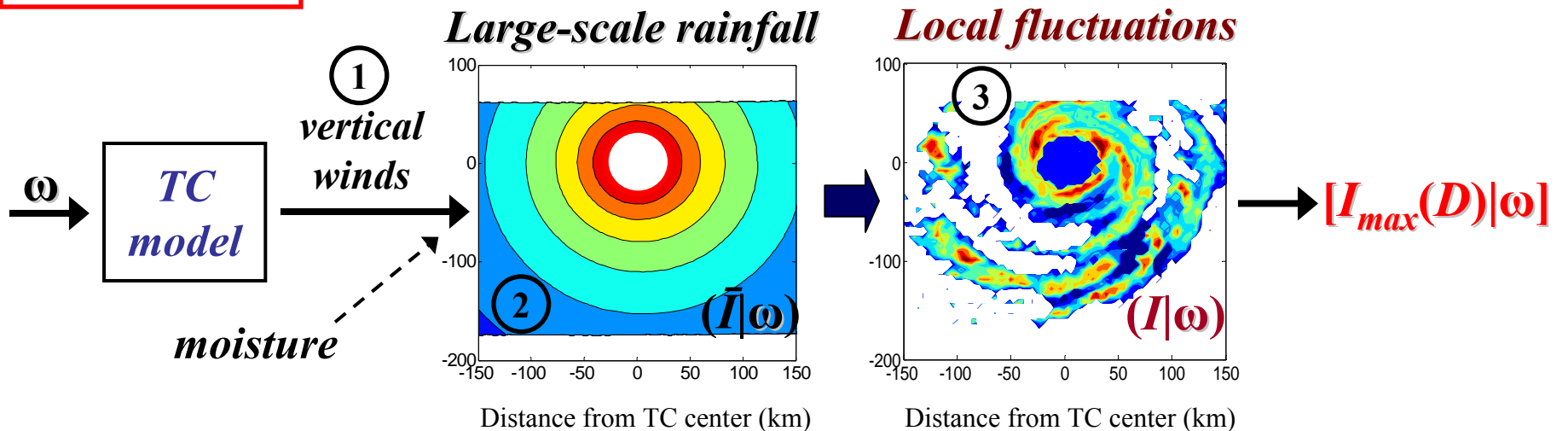
local recurrence (literature)

Implementation

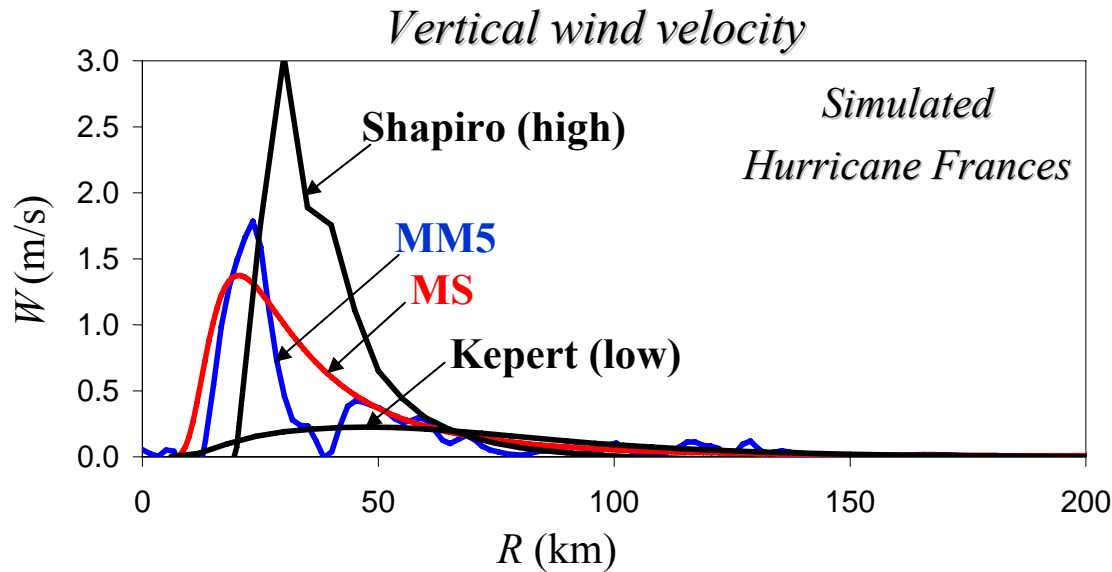
➤ parameters $\omega = [V_{max}, R_{max}, V_t, y]$



$$[I_{max}(D)|\omega]$$



TC Model: Vertical winds and Rain



Vertical winds

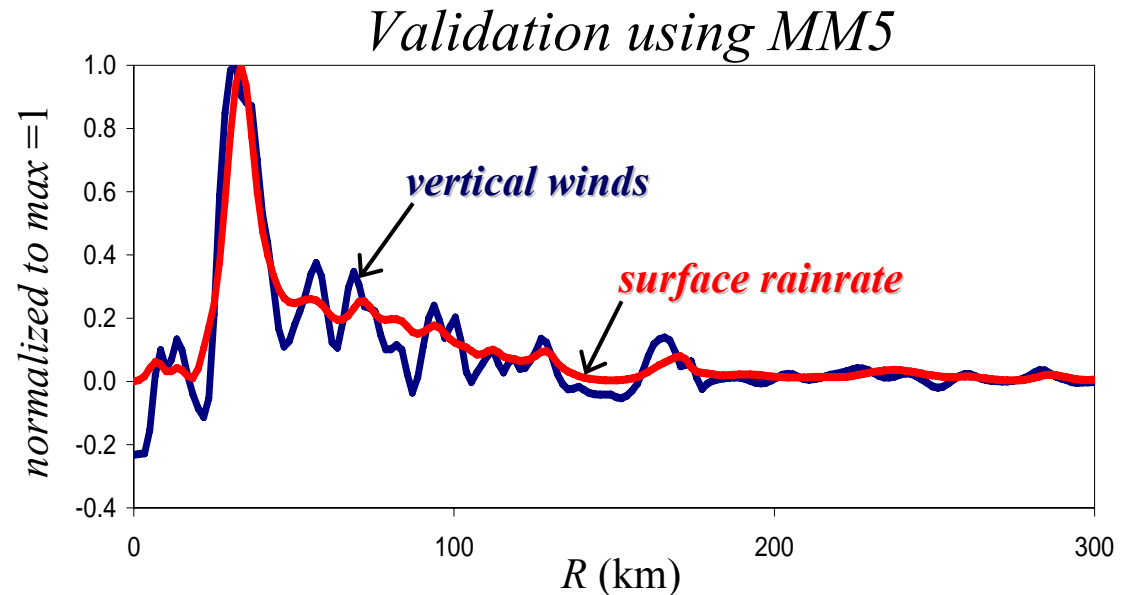
Rain:

$$\bar{I} = c W$$

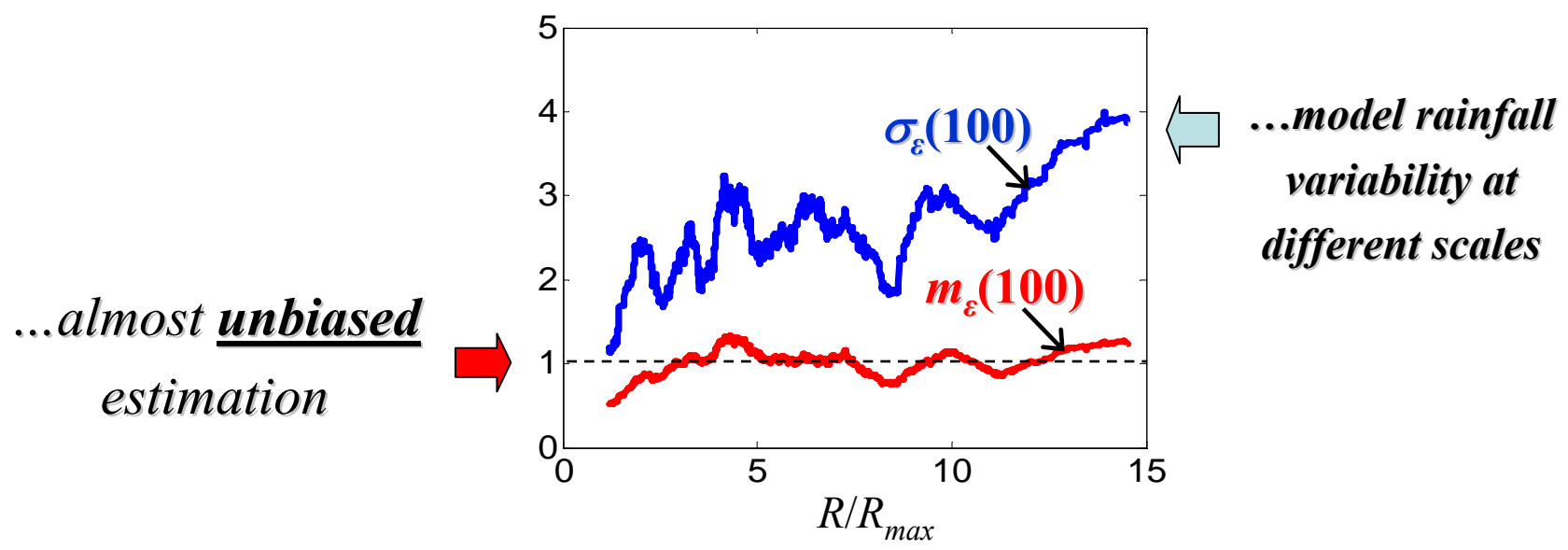
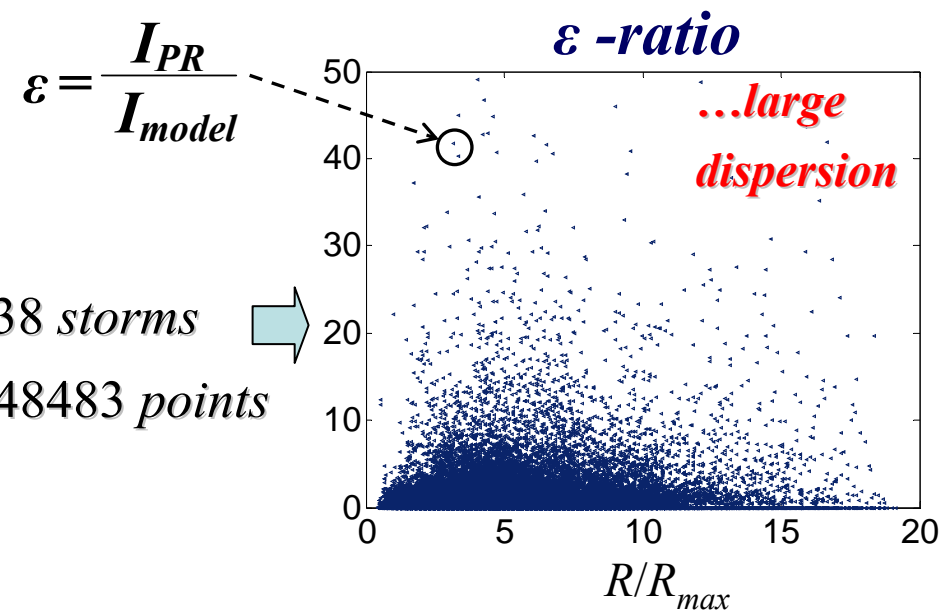
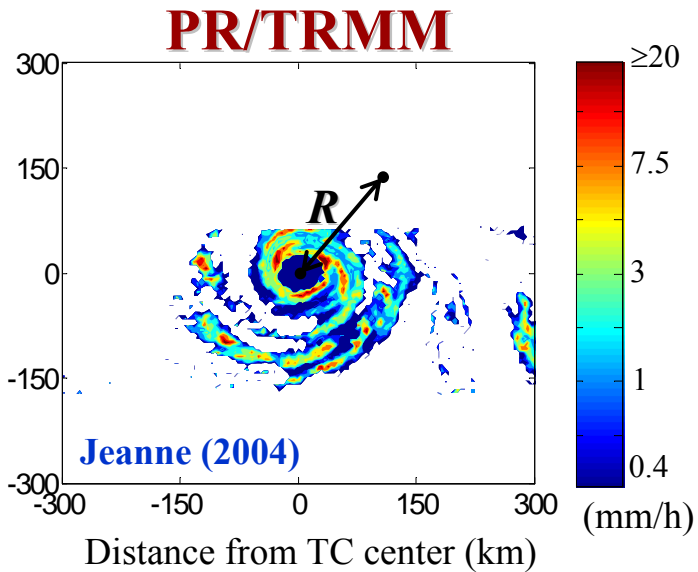
rainrate

moisture content of air

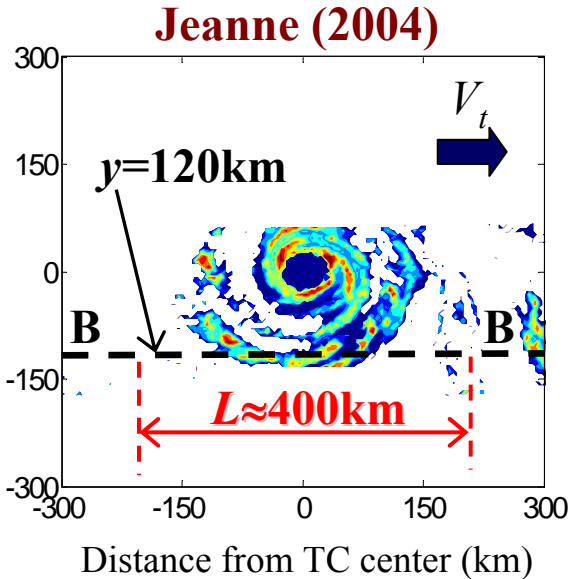
vertical wind speed



Validation using PR/TRMM data

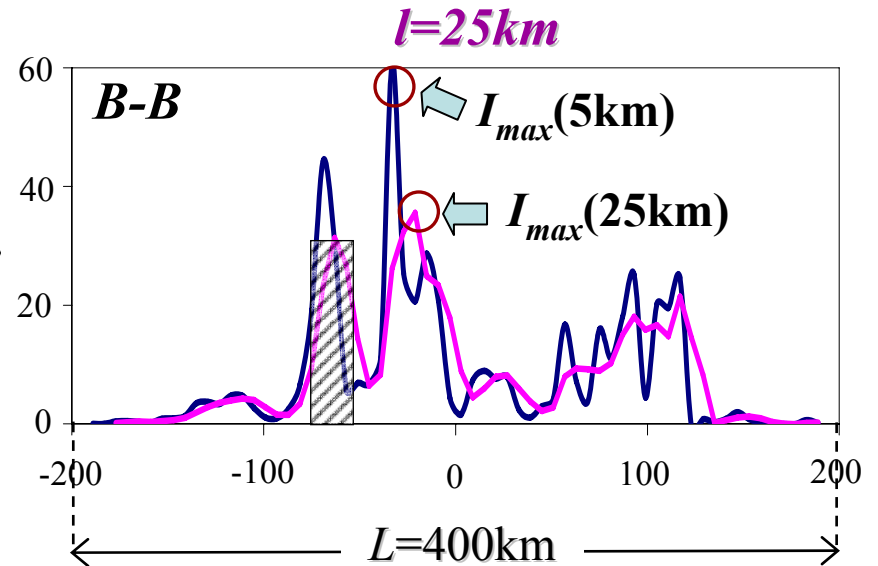
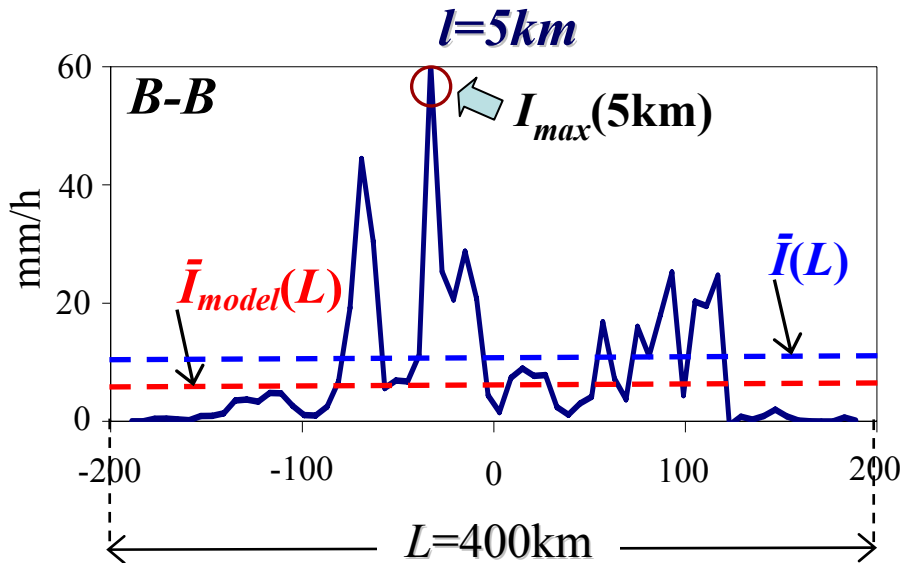


Statistical model of $[I_{max}(l)|\omega]$



$$I_{max}(l) = \overbrace{\bar{I}_{model}(L)}^{(large-scales)} \beta \overbrace{\gamma_{max}(l)}^{(small-scales)}$$

model estimate for the mean rainfall intensity inside L
corrects the model mean relative to the empirical mean
amplification factor for the maximum inside l

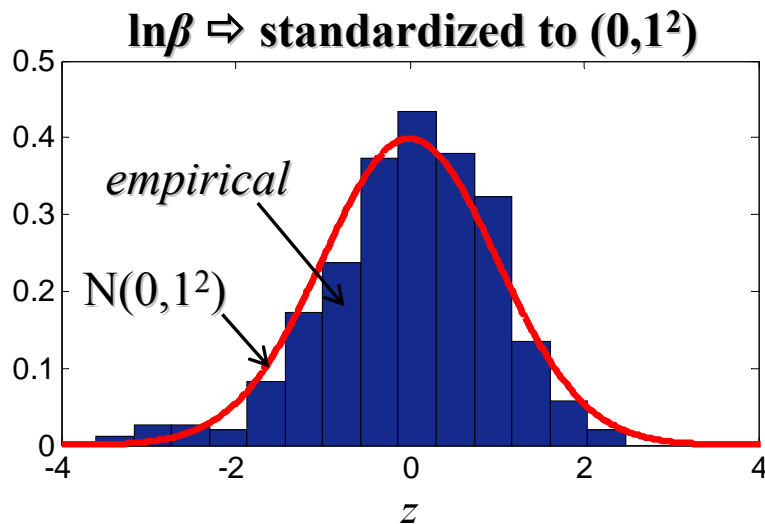


Statistical models for $[\beta|\omega]$ and $[\gamma_{max}(l)|\omega]$

Model for $[\beta|\omega]$

$$\beta = \frac{\bar{I}(L)}{\bar{I}_{model}(L)} \rightarrow \begin{array}{l} \text{empirical} \\ \text{mean inside } L \\ \text{model estimate} \end{array}$$

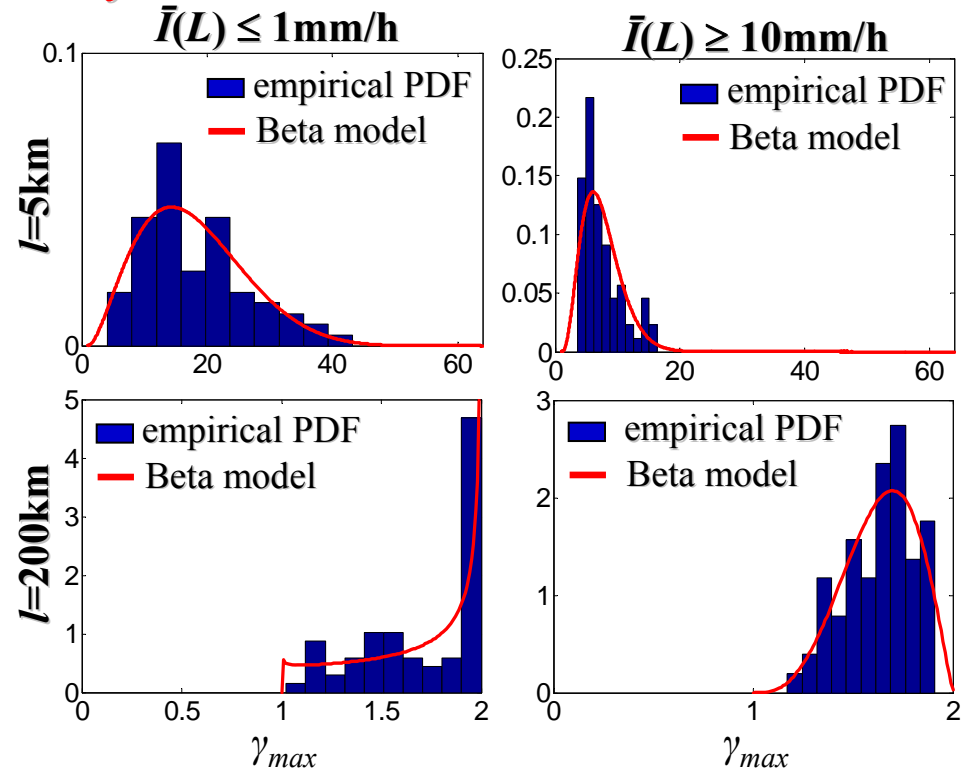
... $\beta(y, \bar{I}_{model}) \sim \text{lognormal}$



Model for $[\gamma_{max}(l)|\omega]$

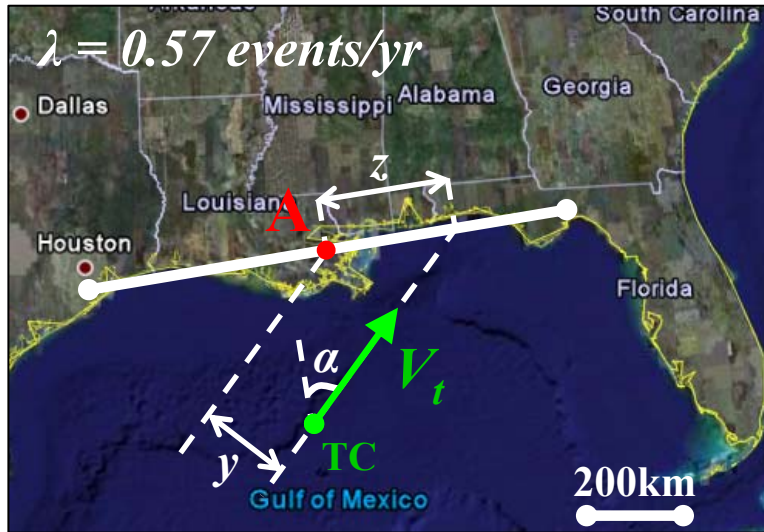
$$\gamma_{max}(l) = \frac{I_{max}(l)}{\bar{I}(L)} \rightarrow \begin{array}{l} \text{maximum rainfall} \\ \text{intensity at scale } l \end{array}$$

parameterize in terms of \bar{I}



Application to New Orleans

➤ Recurrence model for $\omega = [V_{max}, R_{max}, V_t, y]$



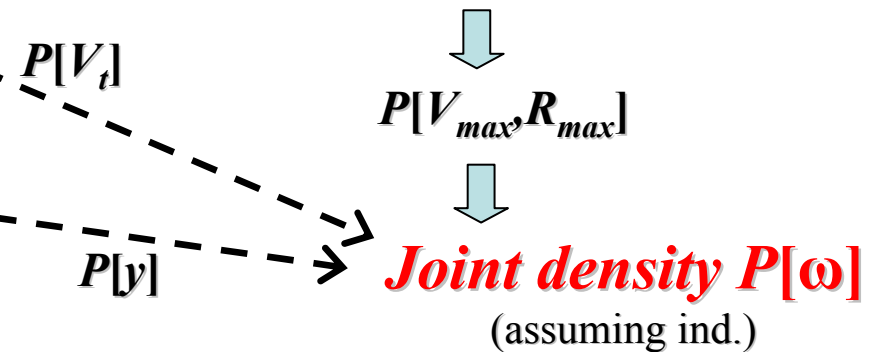
$$\begin{aligned}
 [V_{max}|\Delta P] &\sim \left\{ \begin{array}{l} \text{lognormal with} \\ m = 4.8 \Delta P^{0.559}, \sigma = 0.15 m \\ \text{(Willoughby and Rahn, 2004)} \end{array} \right\} \\
 [R_{max}|\Delta P] &\sim \left\{ \begin{array}{l} \text{lognormal with} \\ m = 3.962 - 0.00567 \Delta P, \sigma = 0.313 \\ \text{(Vickery et al., 2000)} \end{array} \right\} \\
 \Delta P(\text{mb}) &\sim \left\{ \begin{array}{l} \text{shifted lognormal with} \\ m_{\ln \Delta P} = 3.15, \sigma_{\ln \Delta P} = 0.68, \\ \text{Shift par.} = 18\text{mb (IPET, 2006)} \end{array} \right\}
 \end{aligned}$$

(ind.)

$$V_t \sim \left\{ \begin{array}{l} \text{LN with } m = 6\text{m/s} \text{ \& } \sigma = 2.5\text{m/s} \\ \text{(Vickery et al., 2000, Chen et al. 2006)} \end{array} \right.$$

$$\left. \begin{array}{l} z \sim \text{U}[-500\text{km}, 500\text{km}] \\ \alpha \sim \text{N}[-5.4^\circ, (34.9^\circ)^2] \\ \text{(IPET, 2006)} \end{array} \right\} \text{(ind.)}$$

$$y = -z \cos(\alpha)$$

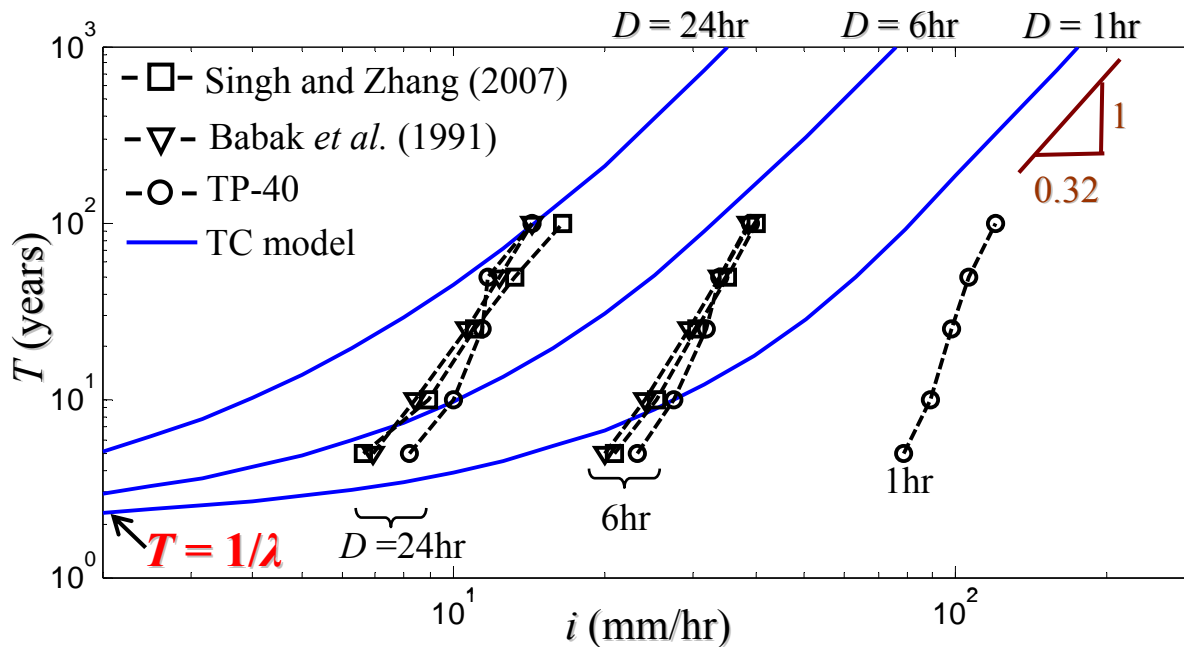
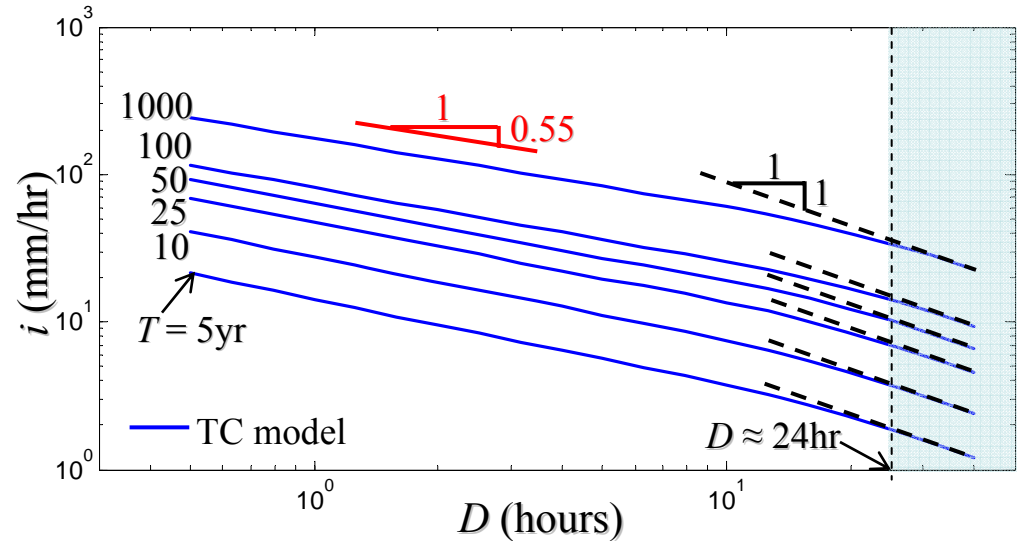


Application to New Orleans: IDF curves

Rainfall Risk and IDF curves:

$$\lambda_D(i) = \lambda \int_{\text{all } \omega} P[I_{\max}(D) > i | \omega] P[\omega] d\omega$$

IDFs: plots of i against D and $T = 1/\lambda_D(i)$ (years)



- For large D and T TCs dominate risk.
- For small D applies the rule: “convection is convection”

Conclusions

➤ We developed a *physical-statistical framework for peak TC rainfall intensities*

- $[I_{max}(D)|\omega]$ {
- *Explicit parameterization* of the hurricane: $\omega = [V_{max}, R_{max}, V_p, \gamma]$
 - *Physical model* to obtain *large-scale rainfall* given ω
(inter-storm variability)
 - *Statistical model* for *rainfall fluctuations*
(intra-storm variability)

➤ *Model validation using PR/TRMM data*

➤ *Application to New Orleans* { *TCs dominate rainfall risk for*
 $D \geq 24h$ and $T \geq 100$ years

Future work {

- *effect of landfall and topography*
- *areal reduction* • *TC wind maxima*